**Atmospheric
Chemistry
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Discussions**

Secondary Organic Aerosol from biogenic VOCs over West Africa during AMMA

G. Capes[1], J. G. Murphy[2,*], C. E. Reeves[2], J. B. McQuaid[3], J. F. Hamilton[4], J. R. Hopkins[4], J. Crosier[1], P. I. Williams[1], and H. Coe[1]

[1]Centre For Atmospheric Science, S. E. A. E. S., University Of Manchester, UK
[2]School of Environmental Sciences, University of East Anglia, Norwich, UK
[3]Institute for Atmospheric and Climate Science, S. E. E., University of Leeds, Leeds, UK
[4]Department of Chemistry, University of York, York, N. Yorkshire, UK
[*]now at: University Of Toronto, Canada

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Abstract

This paper presents measurements of organic aerosols above subtropical West Africa during the wet season using data from the UK Facility for Airborne Atmospheric Measurements (FAAM) aircraft. Measurements of biogenic volatile organic compounds (BVOC) at low altitudes over these subtropical forests were made during the African Monsoon Multidisciplinary Analysis (AMMA) field experiment during July and August 2006 mainly above Benin, Nigeria and Niger. Data from an Aerodyne Quadrupole Aerosol Mass Spectrometer show a median organic aerosol loading of $1.08\,\mu g\,m^{-3}$ over tropical West Africa, which represents the first regionally averaged assessment of organic aerosol mass (OM) in this region during the wet season. This is in good agreement with predictions based on aerosol yields from isoprene and monoterpenes during chamber studies and model predictions based on partitioning schemes, contrasting markedly with the large under representations of OM in similar models when compared with data from mid latitudes.

1 Introduction

Atmospheric aerosols play a key role in climate and atmospheric chemistry. The composition of atmospheric aerosols worldwide is often dominated by the organic fraction (Zhang et al., 2007; Andreae and Crutzen, 1997) and in many cases, the organic aerosol mass is dominated by secondary material (Kanakidou et al., 2005; Volkamer et al., 2006). Measurements of organic aerosol from anthropogenic pollution are widespread in the northern hemisphere, whereas measurements in pristine tropical environments and in particular large areas of Africa remain less well represented in the literature.

Current models based on gas/aerosol partitioning schemes systematically underestimate the measured organic aerosol mass loading in continental northern mid-latitudes (e.g. Jang et al., 2002) and the discrepancy between modelled and measured values increases with photochemical age (Volkamer et al., 2006); this underestimation

highlights our limited understanding of secondary organic aerosol (SOA) formation (and loss) processes. Biogenic volatile organic compounds (BVOCs) are thought to be emitted in greater quantities than anthropogenic VOC emissions on a global basis (Guenther et al., 1995, 2000), so the formation potential for biogenic secondary organic aerosol (BSOA) is substantial (Kavouras et al., 1998), particularly in the tropics where biogenic emissions (Guenther et al., 1995) and concentrations of hydroxyl radicals (OH) (Spivakovsky et al., 1990) are at a maximum. Organic aerosols can exert a direct radiative effect by scattering and absorbing solar and terrestrial radiation but also indirect radiative effects, through their role as cloud condensation nuclei (CCN). Furthermore, predicting human influence on aerosol radiative effects is dependent on quantifying the natural particle system; therefore an inability to predict the quantity and spatial distribution of naturally produced organic particulate matter has important climatic implications.

The UK BAe-146 research aircraft, operated by the Facility for Airborne Atmospheric Measurements (FAAM), was based at Niamey airport in Niger (2.17° E, 13.48° N) between 17 July 2006 and 17 August 2006 to coincide with the onset of the wet season as part of the African Monsoon Multidisciplinary Analysis (AMMA) programme. We show here data from 17 of the 24 sorties flown which give good coverage along a north-south transect between the semi-arid Niamey region of southern Niger and large areas of tropical forest in Benin at low altitudes. During the monsoon season, the prevailing low level (925 hPa) wind direction was south westerly (Janicot et al., 2008), blowing from the Gulf of Guinea over the forested regions of Benin towards Niamey in the north of the operating region. A very clean aerosol background was observed across the region and organic aerosols were detected over forested regions. It should be noted that most of the region is quite rural and there are no major anthropogenic fossil fuel pollution sources. This study compares the regional mass loading with modelling estimates and measured yields from chamber experiments. We present here a regionally averaged mass loading for organic aerosol over West Africa during the 2006 wet season; to our knowledge this is the first time such a measurement has been performed.

2 Aircraft instrumentation

Data was collected on a modified BAe-146 aircraft, which contains a range of instruments to measure aerosol number and composition, trace gas concentrations, and standard meteorological variables.

An Aerodyne Research Inc. Quadrupole Aerosol Mass Spectrometer (Q-AMS) was used to provide near real time mass loadings and chemical composition of the non-refractory components of submicron aerosols. The instrument samples aerosol into a vacuum through an aerodynamic lens, which focuses the particles at a heated vaporizer, where they volatilise. The gas plume is ionised using electron ionisation at 70 eV and the ion fragments are analysed using a quadrupole mass filter and electron multiplier. The Q-AMS has been described thoroughly in previous publications (Jayne et al., 2000; Jimenez et al., 2003; Allan et al., 2003) and the particular aircraft installation of the Q-AMS on the BAe-146 is described in detail by Crosier et al. (2007). The thermal vaporizer is maintained at $\sim$600°C, so the Q-AMS is insensitive to refractory components such as black carbon and mineral dust, which do not vaporize at this temperature. The Q-AMS was calibrated for mass quantification (Ionisation Efficiency, IE) pre- and post-flight using the method of Jimenez et al. (2003) and error estimates are made following the procedure of Allan et al. (2003). The particle beam is blocked every few seconds to measure the signal due to the background in the instrument and this is subtracted from the sample signal. Instrument noise can therefore be estimated by placing a particle filter in the sample inlet. Such estimates show the noise for a 30s sample to be around $1.7\,\mu g\,m^{-3}$ for organics at an altitude of 1000 m (Crosier et al., 2007). All Q-AMS data are reported in mass concentrations at standard temperature and pressure (273.15 K, 101.325 kPa).

Previous studies (e.g. Canagaratna et al., 2007; and studies therein) have shown that when particles are solid a significant number may bounce off the heater, reducing the collection efficiency (CE) below unity. The CE is often evaluated by comparison with other instruments such as PILS-IC or filter measurements and has been shown to be around 0.5 for sulphate aerosol (Drewnick et al., 2005). Particle phase plays a

central role in the CE: liquid particles are sampled with a CE of 1 yet the vast majority of ambient measurements from a range of different environments are well characterised by CEs between 40–50% (Matthew et al., 2008). For pure biogenic SOA, where the particles tend to be liquid aerosols, the CE tends to 100% (Alfarra, 2004). However,

5 sulphate mass loadings over the region were comparable to those of organics, and in previous locations where the sulphate and organics loadings were approximately 50:50 the CE has been shown to be around 0.5 (Zhang et al., 2005). The aerosol during AMMA were bulk neutralized within the uncertainties of the measurements (see method of Zhang et al., 2005) and this tends to be associated with a CE around 0.5

10 (Allan et al., 2004). In this study it was not possible to directly evaluate the true CE by comparison with other instruments, and for the reasons discussed above a CE of 0.5 has been used for the data shown.

Measurements of a suite of volatile organic compounds were made by proton transfer reaction mass spectrometry (PTR-MS; Ionicon Analytik). Specific masses correspond-

15 ing to protonated species of interest were measured for one second every 10–15 s. Zeros were carried out periodically in-flight by diverting the sample air stream through a custom built stainless steel tube filled with a platinum-coated quartz wool catalyst (Shimadzu) heated to 400°C. The response of the PTR-MS was calibrated pre- and post-campaign using a multi-component standard (Apel-Riemer). Due to the high ab-

20 solute humidity encountered during the flights, the protonation of VOCs by both H_3O^+ and the water cluster, $H_3O^+H_2O$, was taken into account. Isoprene is assumed to be the major contribution to the signal at mass to charge ratio (*m/z*) 69, thus the accuracy of this measurement is similar to that quoted for the standard (5%). Both methylvinyl ketone (MVK) and methacrolein (MACR) contribute to the signal at *m/z* 71, so strictly

25 speaking it would be necessary to know the sensitivity and relative contribution of each to calculate the sum of their concentrations. Because we have no alternate measurement of their individual concentrations, the accuracy of the reported sum of their concentrations at *m/z* 71 depends on the similarity of their sensitivities. Individual permeation tubes (Eco-scientific) were used to measure their independent responses, which

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showed them to be within 20% of each other, thus a conservative estimate of the total uncertainty of the reported sum (MVK+MACR) is 25%.

Whole air samples (WAS) were collected in silica coated stainless steel canisters (Thames Restek, UK) for off-line analysis of (C_2-C_8) Volatile Organic Compounds (VOCs). Air samples were analysed using a dual channel gas chromatograph with flame ionisation detectors (Hopkins et al., 2003).

A number of higher molecular weight gas phase volatile organic compounds including monoterpenes were collected using carbon based adsorbent tubes filled with Carbopack B, Carbopack C and Carbosieve SIII (Supelco, Bellefonte, PA, USA). Ambient air was drawn from the portside aircraft air sample pipe which is located towards the front of the aircraft extending beyond the skin boundary layer. A stainless steel metal bellows pump (Metal Bellows, USA) was used to generate a positive pressure system as this ensures that any potential leaks in the sampling train do not compromise the integrity of the sample.

Speciated analysis was accomplished using a GC-TOFMS system (Gas Chromatography coupled to Time Of Flight Mass Spectrometry) comprising an HP 6890 (Agilent Technologies, USA) gas chromatograph and a Pegasus III TOF-MS (LECO, St. Joseph, MI, USA) which is based in the analytical research laboratory at the University of York. The analytical column was a non-polar methyl polysiloxane DB5 type ($60\,m \times 0.25\,mm$ i.d. $\times 1\,\mu m$ film thickness) from J&W Scientific (Folsom, CA, USA). A detailed description of the analytical system and operation parameters can be found in Saxton et al. (2007). Quantification was possible for 4 monoterpenes; α-pinene, limonene, β-pinene and camphene. Several other monoterpenes were observed but without good standards it was not possible to provide quantitative data for them and thus they were not reported here. Unsampled tubes were kept as field blanks to measure contamination during handling. Backgrounds ranged from less than 20 to 30 pptV for the instrument to a maximum of 25 to 150 pptV for an uncapped handling blank. Based on measurements of the gas standards, the precision was estimated to be 22 to 34% (Coefficient of Variation) and accuracy 20%.

3   Results

The measurements made at low altitude over forested areas revealed a strong correlation between VOC concentrations and vegetation density. Isoprene has an atmospheric lifetime of less than 1 h during the daytime in the conditions experienced during AMMA,
5   so its spatial distribution on these scales can be used as an indicator of local vegetative emissions. BVOCs such as monoterpenes and isoprene are known precursors for SOA formation (Claeys et al., 2004; Kavouras et al., 2000) and it is where BVOC concentrations are high that SOA formation is most likely.

Isoprene is emitted in high concentrations from many species of vegetation in tropi-
10   cal regions (Kuhn et al., 2002). During the field project high concentrations of isoprene and its oxidation products methacrolein (MACR) and methyl vinyl ketone (MVK) were found at altitudes below 2000 m over forested regions (e.g. concentrations up to 2 ppb of isoprene were recorded at 500 m altitude – see Fig. 1 for vertical profiles). It is known that different plant types have widely varying emission rates of isoprene and monoter-
15   penes and their relative emission rates can also be highly variable from species to species (Otter et al., 2002; Guenther et al., 1996; Greenberg et al., 2004). Given that this region has a very varied range of plant species (Guenther et al., 1996; Saxton et al., 2007; Greenberg et al., 2004) there will inevitably be a significant amount of spatial averaging of the emission profiles from individual species when measured from
20   an airborne platform. The mean α-pinene concentration at altitude <2000 m over the forested regions was 38 ppt; the mean for isoprene was 610 ppt. The α-pinene concentration will be an overestimate relative to isoprene because the tube samples were collected specifically when high VOC concentrations were expected, compared to the continuous sampling of the PTR-MS and Q-AMS.
25   Typical aerosol number concentrations in the region were low (particle number concentration greater than 3 nm<500 cm^{-3}) which indicates a very clean background (Fig. 2). Number concentrations shown have been screened to remove influence from other sources such as biomass burning and large urban centres. Biomass burning

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influence was identified by enhancement in acetonitrile concentration, whereas the only major urban plume sampled by the aircraft (Lagos) was identified geographically and by enhanced CO concentrations. The data were further screened to investigate the change in number concentration in air sampled with high or low concentrations of isoprene, using a threshold of 100 ppt (subsequently referred to as the I_{high} and I_{low} subsets). The I_{high} screening criteria used are intended to reveal the biogenic signature over the forest, whereas the I_{low} air represents highly aged regional air. Typical number concentrations for the I_{high} and I_{low} subsets are very similar, less than $500\,cm^{-3}$ in each case.

In forested regions in mid and high latitudes, an order of magnitude enhancement in the number concentration over background values has been observed during nucleation events with particle concentrations around $10^4\,cm^{-3}$ (Kulmala et al., 2001); this was not evident during AMMA as can be seen in Fig. 2. This lack of new particle formation is in agreement with that previously observed from a ground station in the Amazon (Zhou et al., 2002; Rissler et al., 2004).

Low loadings of aerosol particle mass were observed throughout the whole region when the data were screened to remove the occasional influences of biomass burning and urban pollution. For this reason, mass loadings during AMMA have been determined on the basis of statistical comparison. Highest concentrations of BVOCs were observed at altitudes below 2000 m (see Fig. 1) and the Q-AMS has improved signal-to-noise when the sample pressure is high (Crosier et al., 2007), so only data points below 2000 m were selected for this estimate. Figure 3 shows organic mass frequency distributions for I_{high} and I_{low} periods; the mean, lower quartile, median , upper quartile and standard deviations of these distributions are summarised in Table 1. This analysis shows an enhancement in OM for the air with elevated isoprene concentrations, which would tend to indicate BSOA formation, though the enhancement is small.

Many of the data points in the I_{high} subset also exhibited elevated mean benzene concentrations (96 ppt compared with 76 ppt for the I_{low} data), which is indicative of anthropogenic pollution sources e.g. Fenger (1999). Benzene is a much longer lived

VOC than isoprene and so may have been transported considerable distance from its source. It is therefore possible that the observed OM has a contribution from dilute anthropogenic pollution – cf. benzene concentrations of 1–2 ppb in polluted environments (Rappengluck et al., 2000; Steinbacher et al., 2005). However, it has been observed that toluene is emitted from some vegetation (White et al., 2009) and it may be that similar emissions contribute to the small enhancement in benzene observed. Sulphate aerosols were observed in the I_{high} subset i.e. in the same region as enhanced OM (see Fig. 4). However, similar concentrations of sulphate were present throughout the whole dataset and moreover the presence of sulphate is inconclusive as an indicator of anthropogenic influence. A major source of sulphate aerosols is SO_2 from anthropogenic activity, but there are also biogenic sources, which may be significant in this region; the observed mass concentrations are consistent with aerosol formation from vegetative emissions of dimethylsulphide (DMS) and H_2S (Andreae et al., 1990).

With the notable exception of isoprene, the WAS data show low VOC concentrations in both the I_{high} and I_{low} air, compared with regions containing large anthropogenic sources e.g. typical rural UK values are a factor of 2–10 higher (Fowler et al., 1997); species such as ethane and propane dominate the VOC signature over West Africa which indicates very aged air. This is consistent with the very low NO_x concentrations over the region (median 0.21 ppb, inter-quartile range 0.31 ppb for I_{high}). Hence, while it is most likely that the OM in this study is dominated by the biogenic contribution, some contribution from anthropogenic sources cannot be entirely ruled out. These data therefore represent an upper limit to the biogenically produced SOA loading.

4 Theoretical SOA estimates for West Africa

Isoprene, monoterpenes and sequiterpenes are thought to be the major SOA precursors globally (Kroll et al., 2006; Kanakidou et al., 2005) so an estimate for SOA formation should be based on these compound classes. Isoprene and monoterpenes have been shown to be emitted from forested regions of the African Sahel in high concentrations during the wet season (Saxton et al., 2007), but no quantitative measurements

for sesquiterpenes are available. Our estimate is therefore based on isoprene and monoterpenes only and hence is likely to be an under prediction. A simple approach to estimating the SOA formation from the observed isoprene and monoterpenes is presented using data from the regions defined by the I_{high} screening criteria outlined for Fig. 3. Isoprene, α-pinene, ß-pinene and limonene were the precursor species measured for which relevant emission ratios were readily available.

Chamber estimates of isoprene derived SOA suggest mass yields around 3% (Kroll et al., 2006), and molar yields of MVK and MACR from isoprene oxidation have been shown to be 15% and 18% respectively in low NO_x conditions (Ruppert and Becker, 2000). The lifetime of MVK+MACR is around 8 h (Atkinson et al., 2006), so there is decay on the time scale of a day. The SOA precursor may exist for longer, but chamber experiments show that the bulk of the aerosol mass is formed on timescales of several hours (Kroll et al., 2006; Alfarra et al., 2006) i.e. the timescale over which the organic aerosol is produced is similar to the timescale for MVK + MACR oxidation. Assuming that the only source of MVK and MACR is isoprene degradation these isoprene oxidation products should be present in the ratio of their yields. Converting from molar to mass yields and mixing ratios to mass concentrations yields an estimate for SOA given by Eq. (1).

$$SOA_{Iso} = [MVK + MACR] \times \frac{Y_{SOA}}{Y_{MVK}\left(\frac{M_{MVK}}{M_{Iso}}\right) + Y_{MACR}\left(\frac{M_{MACR}}{M_{Iso}}\right)} \times n \times \frac{mol \times M_{MVK+MACR}}{N_A} \quad (1)$$

where SOA_{Iso}=SOA mass concentration from isoprene oxidation, [MVK + MACR]=concentration of the sum of MVK and MACR (ppt), Y_{SOA}=aerosol mass yield=0.03 (Kroll et al., 2006), Y_{MVK}=molar yield of MVK=0.15 (Ruppert and Becker, 2000), Y_{MACR}=molar yield of MACR=0.18 (Ruppert and Becker, 2000), n=number of molecules in 1 m^3 air=2.5×10^{25}, mol=number of moles=1, N_A=Avogadro's number=6.02×10^{23}, M_{Iso}=molar mass of isoprene=68 gmol^{-1}, M_{MVK}=molar mass of MVK=70 gmol^{-1}, M_{MACR}=molar mass of MACR=70 gmol^{-1}, $M_{MVK+MACR}$=molar mass of MVK and MACR=70 gmol^{-1}. The mean MVK+MACR

concentration for the I_{high} subset was observed to be 430 ppt. Equation (1) gives a prediction of organic mass loading from isoprene degradation of $0.114\pm0.063\,\mu g\,m^{-3}$ over the region. The uncertainty quoted is the standard deviation and reflects variability in the MVK+MACR concentration; the choice of Y_{SOA} reflects the upper limit of isoprene-derived SOA. This estimate assumes complete oxidation of isoprene to achieve the final concentration of SOA. Certainly, the methylvinyl ketone and methacrolein concentrations in Figure 1 are similar in magnitude to that of isoprene, indicating that considerable oxidation has taken place and the mixing time in the atmosphere is of the order of several hours. Whilst chamber measurements show that SOA continues to form for up to 24 h after the start of oxidation, much of the aerosol mass has been formed in the first half of these studies. Hence, though our calculation may underestimate the maximum potential SOA from isoprene oxidation, it is not likely to be a very large under-prediction.

An estimate of the contribution of monoterpenes to the SOA can be derived by extending the approach above. Monoterpenes, like isoprene are short lived in the atmosphere whereas aerosols are long lived, and hence their contribution to the measured SOA is due to the integrated amount of monoterpene emitted into the air mass and is nearly independent of the in situ abundance of the BVOC. Furthermore the 1/e lifetimes of α-pinene and isoprene under conditions experienced during AMMA were both around 2.1 h. Hence we assume that the total quantity of a particular VOC that has reacted to form products (denoted by the subscript TR) arises from the total quantity of the VOC emitted (denoted by the subscript TE) see Eq. (2). Using an α-pinene (AP)/isoprene (Iso) emission ratio of 0.5 ± 0.05 to represent AP_{TE}/Iso_{TE} (derived from measurements made in Benin during the same time period as our measurements (Saxton et al., 2007)), and using a suitable aerosol yield from α-pinene, $Y_{AP}=0.15$ (Ng et al., 2006), we arrive at an estimate for SOA from α-pinene expressed by Eq. (3), $SOA_{AP}=0.29\pm0.19\,\mu g\,m^{-3}$, though this figure may represent the lower limit as yields up to 40% have been reported (Ng et al., 2007).

$$AP_{TR} = Iso_{TR} \times \frac{AP_{TE}}{Iso_{TE}} \qquad (2)$$

$$SOA_{AP} = \frac{SOA_{Iso}}{Y_{Iso}} \times \frac{AP_{TE}}{Iso_{TE}} \times Y_{AP} \qquad (3)$$

The same approach was taken for β-pinene (BP) and limonene using emission ratios BP/Iso=0.5±0.05, Limonene/Iso=1±0.1 (Saxton et al., 2007) and yields Y_{BP}=0.032, $Y_{Limonene}$=0.087 (Griffin et al., 1999a) to give SOA$_{BP}$=0.061±0.040 μg m^{-3}, and SOA$_{limonene}$=0.33±0.21 μg m^{-3}. Uncertainties in these estimates reflect the uncertainty in the emission ratios used, in addition to the uncertainty in SOA$_{Iso}$. For comparison, a range of SOA estimates have been generated in this manner using emissions data from various sources. These reflect the high degree of variability in emission profiles between different locations with similar biome and are summarized in Table 2. However, measurements taken in Amazonia may not be representative of BVOC in West Africa, and so the Saxton et al. (2007) data were chosen as those most pertinent to this study.

Clearly our estimate based on four precursor species only accounts for a fraction of the contribution to the total SOA formation potential from BVOC e.g. α-pinene, β-pinene and limonene represent 81% of monoterpene emissions on a global basis (Griffin et al., 1999b). Given that the monoterpene measurements by PTR-MS were mostly below detection limit, this contribution cannot be more accurately determined. Our calculations predict a total SOA concentration of 0.795±0.503 μg m^{-3} compared with the median observed value of 1.08±0.06 μg m^{-3}. The measured OM will inevitably contain a small contribution from pre-existing OM onto which the SOA partitions; an unknown contribution which is therefore not included in the calculation. Nevertheless, agreement to within the uncertainties based on a few major BSOA precursors suggests that BSOA formation can readily explain the observed OM. In fact, our calculations are likely to also under-represent the BSOA from isoprene and α- and β-pinene for the reasons discussed above.

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5 Discussion

Organic aerosol mass loadings measured from an aircraft during the programme provide the first regional measurements of organic particulate mass over clean forested regions of sub-tropical West Africa during the growing season. A median concentration of organic submicron aerosol mass 1.08 μg m^{-3} was observed. The measured values are comparable to previous measurements taken in tropical environments (e.g. 2.93 μg m^{-3} in Amazonia (Artaxo et al., 2002)). Although the measurements from the African Sahel during the wet season reported in this paper are somewhat lower than those reported from Amazonia, this is reflected in lower VOC concentrations than typical for Amazonia (2–5 ppb isoprene and more than 100 ppt α-pinene have been reported (Karl et al., 2007; Greenberg et al., 2004)).

Global model estimates of OM in this region of West Africa derived from bulk yields and partitioning schemes cover the range 0 μg m^{-3}–3 μg m^{-3} (Chung and Seinfeld, 2002; Tsigaridis and Kanakidou, 2003; Lack et al., 2004). The Q-AMS derived regional organic mass loadings fall within the range of estimates from these model simulations and are consistent within the large uncertainties associated with partitioning schemes in models and with our bulk yield calculations using the BAe-146 VOC data. Furthermore, our simple estimates of SOA formation shown in Table 2 of between 0.2 and 0.8 μg m^{-3} are consistent with both the observations and the global model results. This marks a strong contrast with urban environments where secondary material often dominates the organic mass fraction e.g. (Crosier et al., 2007; Turpin and Huntzicker, 1995) and is far in excess of that predicted by model simulations by between one and two orders of magnitude e.g. Volkamer et al. (2006). Moreover, there is some evidence to suggest an anthropogenic contribution to the OM measured, our estimates of SOA did not include any contribution from sesquiterpenes and our budgets are based on the oxidation of isoprene and the monoterpenes being completed. All this suggests that models are possibly over predicting in this environment. Tsigaridis and Kanakidou (2003) find their simulations consistently underestimate OM in mid-latitude rural environments compared with field measurements from a number of locations around

the globe; similarly Chung and Seinfeld (2002) found consistent under prediction of OM compared with measured values in rural areas of the USA.

The reasons why models under predict OM in the presence of anthropogenic pollution, yet not for clean tropical environments, remain unclear. Elemental carbon/ organic carbon (EC/OC) measurements using ^{14}C have shown that 60%–70% of OM is modern carbon i.e. biogenic in nature, in studies in Europe (Szidat et al., 2004) and in the US (Lemire et al., 2002). These studies were performed in locations where large excesses of OM have been reported (Volkamer et al., 2006) and as such differ from this work. However, they are in regions where significant anthropogenic pollution occurs; this is not the case in West Africa which has far fewer anthropogenic pollution sources. Anthropogenic pollution may influence the partitioning of organic matter to the particle phase by enhancing oxidation rates – increasing photochemical processing – enhancing OH and O_3; by changing product channels due to the presence of NO_x; by also changing the nature of the particulate by the addition of significant inorganic salts that are hygroscopic and/or acidic in nature. The reasons for the differences in levels of agreement between measurement and model prediction of organic particulate between the polluted mid-latitudes and tropical forests remain unclear. Better elucidation of the behaviour of organic carbon chemistry and its condensation to the particle phase is clearly necessary if predictive models of OM are to be effective over a broad range of environments.

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Table 1. Organic mass loadings (μg m^{-3}) for data screened for various criteria. I_{high} and I_{low} indicate data below 2000 m (excluding urban centres, biomass burning influences and data outside the latitude range 7–13° N), selected for isoprene above and below 2000 m. Q1 and Q3 are the lower and upper quartiles. The mean is quoted ± standard error.

Criteria	Mean ± std err	Median	Q1	Q3	Standard Deviation
All data	0.216±0.096	0.0310	−1.786	1.564	6.36
All <2000 m	1.18±0.16	0.614	−0.642	1.90	8.06
I_{low}	0.520±0.112	0.480	−0.696	1.626	1.804
I_{high}	1.164±0.062	1.072	0.125	2.14	1.762

Table 2. SOA estimates (μg m^{-3}) from Eq. (3) using a range of emission ratios. The tubes data refers to ambient concentrations from the adsorbent tubes, but isoprene and α-pinene have very similar 1/e lifetimes under the conditions during AMMA (2.10 h and 2.08 h, respectively) and so the ratio of their concentrations should remain approximately constant from source. SOA$_{total}$ includes the contribution from calculated SOA$_{Iso}$.

	Saxton et al., 2007	Tubes data	Greenberg et al., 2004	Kesselmeier et al., 2000
SOA$_{\alpha p}$	0.29±0.19	0.034±0.037	0.063±0.043	0.032±0.016
SOA$_{\beta p}$	0.061±0.040	n/a	n/a	0.052±0.033
SOA$_{limonene}$	0.33±0.21	n/a	n/a	0.016±0.018
AP/iso	0.5	0.06	0.11±0.029	0.57±0.28
BP/iso	0.5	n/a	n/a	0.17±0.09
Limonene/iso	1	n/a	n/a	0.11 ± 0.55
SOA$_{total}$	0.795±0.503	0.172±0.3	0.177±0.106	0.214±0.13
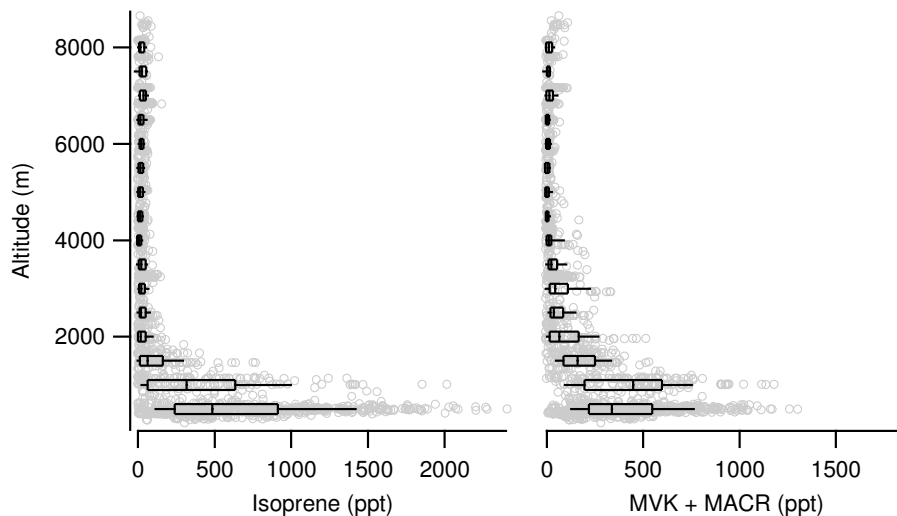
Fig. 1. Vertical profiles show isoprene and its degradation products methyl vinyl ketone (MVK) and methacrolein (MACR) between 7–13° N. Grey markers are individual data points. The box and whiskers show concentrations at 500 m intervals, at the mid point of each 500 m altitude bin. The box centre denotes the median concentration; box edges denote the upper and lower quartiles; whisker extremities denote the 10th and 90th percentile concentrations.
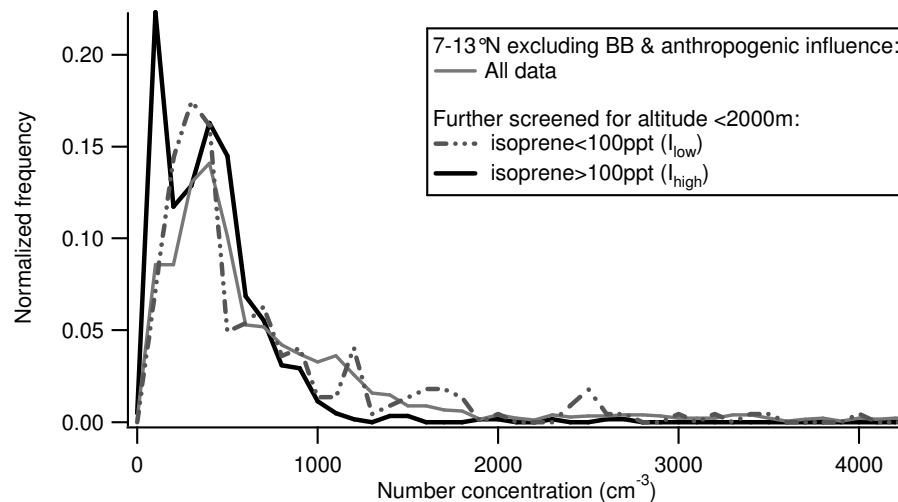
Fig. 2. Frequency distribution for particle number concentration, shown for all data between 7 and 13° N excluding periods of biomass burning and anthropogenic influences and data further screened according to altitude and isoprene concentrations.
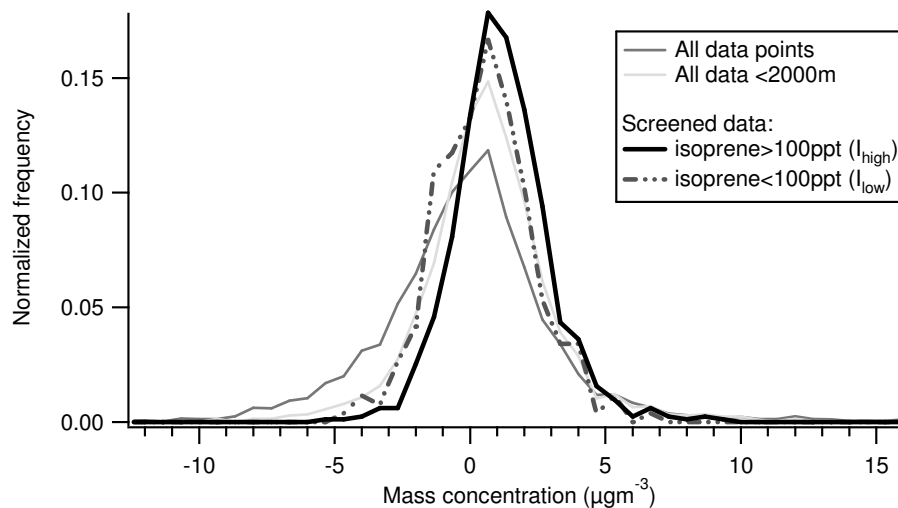
Fig. 3. Frequency distributions for organic aerosol mass. The screened data is below 2000 m and excludes urban centres and biomass burning influences and data outside the latitude range 7–13° N.
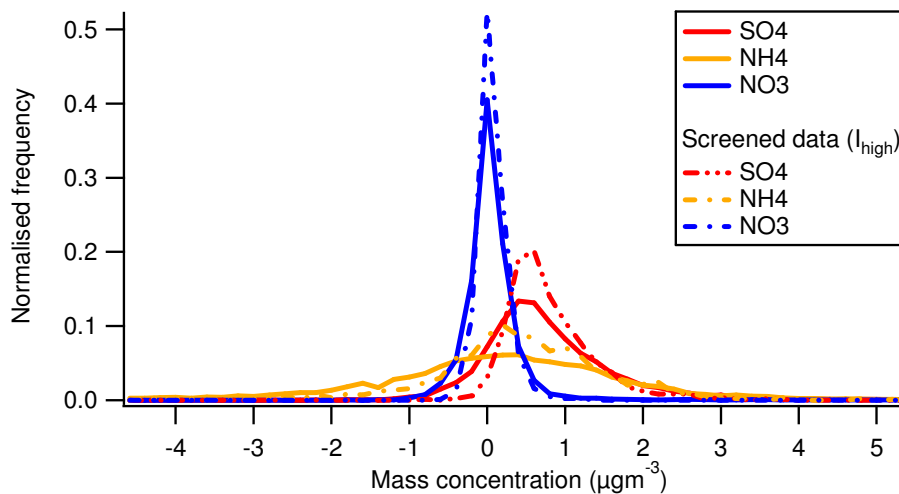
Fig. 4. Frequency distributions for inorganic aerosol mass. "Screened" data refers to data below 2000 m with isoprene >100 ppt, excluding urban centres, biomass burning influences and data outside the latitude range 7–13° N.