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**Temporal trends of anthropogenic SO<sub>2</sub> emitted by smelters**

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# Temporal trends of anthropogenic SO<sub>2</sub> emitted by non-ferrous metal smelters in Peru and Russia estimated from Satellite observations

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## Abstract

We report on satellite observations of atmospheric Sulfur Dioxide (SO<sub>2</sub>) emitted from metal smelting industries in Peru, South America and Siberia, Russia. Most of the non-ferrous metal ores are sulfidic and during the smelting process the sulfur is emitted as SO<sub>2</sub>. In addition to Norilsk, Russia, Peruvian copper smelters are among the most polluting point sources in the world. We retrieve SO<sub>2</sub> column amounts from spectra of the Global Ozone Monitoring Experiment (GOME) on the Earth Research Satellite 2 (ERS-2) for the years 1996 to 2002 using an algorithm based on differential Optical Absorption Spectroscopy (DOAS). Areas of enhanced SO<sub>2</sub> column amounts are clearly identified on a 7-years mean map of GOME observations over the regions with La Oroya and Ilo copper smelters of Peru and Norilsk smelters of Russia.

Since the instrument sensitivity is highly dependent on surface albedo, SO<sub>2</sub> vertical profile, solar zenith angle (SZA), wavelength, clouds, and aerosol, radiative transfer modelling is used to convert the analysed slant column densities into vertical column densities. In this study, the full spherical Monte-Carlo radiative transport model TRACY-II is used for SO<sub>2</sub> AMF calculation.

GOME data is analysed in further detail by calculating time series over these regions. For the different locations, the results demonstrate both, increasing and decreasing trends in the SO<sub>2</sub> column amounts over the time period of 1996–2002. The decreasing trend for the Ilo copper smelter is in good agreement with implemented measures for emission reductions. However, even for the cases with decreasing trends, these point sources are still a dominant source of anthropogenic SO<sub>2</sub> emissions in their region. For the smelters in Peru, the potential influence due to SO<sub>2</sub> emission by the nearby volcanoes is investigated and found to be negligible.

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## 1 Introduction

Sulfur is prevalent in metal ores that contain common metals like aluminium, copper, zinc, lead, nickel, iron etc. For instance, copper is found in the earth's crust primarily as chalcopyrite ( $\text{CuFeS}_2$ ), with other sulphide ores, such as bornite ( $\text{Cu}_5\text{FeS}_4$ ), covellite ( $\text{CuS}$ ) and chalcocite ( $\text{Cu}_2\text{S}$ ), and oxide ores ( $\text{CuO}$ ) (EET, 1999). The reduction of the metal ion to the free metal is normally accomplished in a process referred to as smelting. As a result sulfur in the metal ore is eliminated as sulfur dioxide ( $\text{SO}_2$ ).

Sulfur dioxide emissions occur in significant amounts at various processes associated with metal smelting. Its release into the atmosphere occurs primarily during material handling operations (Fugitive Emissions), and during smelting processes gases are exhausted through stack (Point Source or Stack Emissions). Recovery of  $\text{SO}_2$  – usually the hot gases are piped to an acid plant for the manufacturing of sulfuric acid before exhausting them through a stack (EET, 1999) – is a good environmental practice (adopted in several industrialized countries) and may be economically beneficial.

Global  $\text{SO}_2$  emissions are mainly of anthropogenic origin and contribute about 76 Mt S/yr, mainly from fossil fuel consumptions and industrial activities with 68 Mt S/yr from Northern Hemisphere and 8 Mt S/yr from Southern Hemisphere (Benkovitz et al., 1996).  $\text{SO}_2$  emitted from smelters has severe effects on forest, vegetation, land surface, and water reservoirs in the close vicinity of the smelters (EET, 1999; Aamlid, 2002; AMAP, 2006; Savard et al., 2006; Telmer et al., 2006). Especially; forests near large sources are severely damaged. Symptoms are recognized on Scots pine (*Pinus sylvestris*) and downy birch (*Betula pubescens*) and on other plants, e.g. dwarf birch (*Betula nana*) and Bilberry (*Vaccinium myrtillus*) (Aamlid, 2002; AMAP, 2006).

Telmer et al. (2006) stated during the study of 99 lakes around the Horne smelter (See Table 1) in Québec that lakes within 50 km of the smelter have elevated metal concentrations in their near-surface sediments due to stack emissions. Therefore, the study of emissions in the environment around smelters is a priority in evaluating policies in relation to a sustainable development.

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Satellites provide the best platform to monitor environmental hazards on global scales over extended periods of time. GOME observations during 1996–2002 showed significant amount of anthropogenic SO<sub>2</sub> around the globe mainly from industrial regions such as USA, China, Eastern Europe, South Africa, Canada, Chile, Japan, Peru, and particularly from Siberia, Russia (Khokhar et al., 2004; Khokhar, 2006).

A database presented in Table 1 shows the information about Cu smelters (selection mainly based on their production capacities) from mostly polluted and industrialised region around the world. This work will focus on SO<sub>2</sub> emissions from Peruvian Cu smelters La Oroya, Ilo and from Norilsk smelters in Siberia, Russia.

## 2 Data analysis

Over the last decades, the potential of remote sensing measurements on satellite platforms was increasingly used to analyze various parameters of the state of the earth, e.g. daily weather, magnetic field, topography, land cover, ocean waves or temperature. Satellite data has also successfully been used to analyze the chemical composition of the atmosphere by measuring spectra of the light reflected (or emitted) by the earth's surface and atmosphere. Such measurements are performed by instruments using different spectral ranges (e.g. UV, visible, IR, and microwave) and viewing geometries, with different potentials and limitations. Additionally, satellite instruments have the advantage of being able to observe the whole globe with the same instrument on a long timescale. Thus, it is possible to gain important details of long-term temporal evolutions and global variability of the atmospheric state.

In 1995, the Global Ozone Monitoring Experiment (GOME), a nadir-scanning UV-VIS spectrometer, was launched on ERS-2 into a near-polar sun synchronous orbit at a mean altitude of 790 km with a local equator crossing time at 10:30 a.m. (Burrows et al., 1999). Total ground coverage is obtained within 3 days at the equator by a 960 km across-track swath. For the measurements presented in this work, the ground pixel size is 40 (along track) by 320 (across track) km<sup>2</sup>. As suggested by its name, GOME was

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originally designed for the monitoring of (mainly stratospheric) ozone. However, the instrument proved to be extremely useful for the determination of many tropospheric trace gases, like e.g. H<sub>2</sub>O, BrO, CH<sub>2</sub>O, NO<sub>2</sub> and SO<sub>2</sub> (e.g. Platt et al., 1994; Wagner et al., 2001 etc.).

5 GOME measures the direct solar spectrum (irradiance) and the Earthshine radiance, i.e. the sunlight reflected by the earth's surface or scattered back by molecules and aerosols in the atmosphere (Burrows et al., 1999). From the ratio of earthshine radiance and solar irradiance measurements, slant column densities (SCD) of the respective absorbers can be derived by applying the technique of differential optical absorption spectroscopy (DOAS) (Perner and Platt, 1979; Platt et al., 1994; Richter et al., 10 1998a, b; Wagner et al., 2001, 2008). For the SO<sub>2</sub> evaluation we made use of the spectral range from about 312 nm to 327 nm where the strongest bands of the A<sup>1</sup>B<sub>2</sub>←X<sup>1</sup>A<sub>1</sub> transition are located. The SCD is the integrated concentration along the light path. For the conversion of the measured SO<sub>2</sub> SCD into the vertical column density (VCD: 15 concentrations integrated vertically) usually an air mass factor (AMF) is applied, which is the ratio of the SCD and the VCD. The AMF can be determined from the numerical simulation of the atmospheric radiative transfer. Also so called Box-AMF can be calculated as the ratios of partial SCD and VCD for different altitude layers. They describe the altitude dependence of the sensitivity of the satellite observation.

### 20 3 Radiative transfer modelling

In case of SO<sub>2</sub>, the calculation of an appropriate AMF is complicated by several factors: First the SO<sub>2</sub> AMF depends strongly on the SO<sub>2</sub> profile, which largely differs for the different types of sources like e.g. fugitive emissions, industrial plants, or Volcanoes (Richter et al., 2002a; Khokhar et al., 2005; Khokhar, 2006). Furthermore, 25 the AMF also strongly depends on the surface albedo, clouds, SZA and aerosol loading. Information about all these parameters is usually not available for a given GOME SO<sub>2</sub> analysis (Khokhar, 2006). Even in the absence of aerosol as result of the strong

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Rayleigh scattering in the UV, the measurement sensitivity decreases strongly towards the surface over dark scenes (i.e. at low surface albedo). A sensitivity study on SO<sub>2</sub> AMF calculations presented in (Khokhar et al., 2005) proposed to apply an average AMF of unity. This means that on average the SO<sub>2</sub> VCD actually equals the retrieved SCD. This simplification can be useful e.g. for trend analysis, but in individual cases large deviations from an AMF of about unity can occur, depending especially on cloud cover and surface albedo. Below, we present detailed radiative transfer simulations to exploit the range of possible AMFs relevant for this study.

The calculation of appropriate AMFs for satellite observations of tropospheric SO<sub>2</sub> was investigated in several studies, for instance Thomas et al. (2004), Richter et al. (2002a) and Khokhar et al. (2005, 2006). Here, the full spherical Monte-Carlo model TRACY-II model is used for SO<sub>2</sub> AMF calculations (Deutschmann and Wagner, 2007; Wagner et al., 2007). To account for all aforementioned effects several specific cases were modelled. Sample results presented in Fig. 1 (right panel) show that the SO<sub>2</sub> AMF (for the vertical profile shown in Fig. 1 (left panel) (Taubman et al., 2006) strongly depends on surface albedo, SZA, and the presence of aerosols. The instrument sensitivity towards tropospheric SO<sub>2</sub> expressed as Box-AMF, for different surface albedi is given by Fig. 1 (left panel). The calculations are performed for a wavelength of 315 nm, a SZA of 40°, an aerosol extinction profile as shown in Fig. 1 (see Frieß et al., 2005), single scattering albedo (SAA) of 0.95, an asymmetry parameter  $g=0.68$ , surface albedo of 5% (dark surface, solid gray line) and 50% (bright surface, gray line with crosses). As shown in Fig. (1), for bright surfaces (e.g. over snow and ice) much higher sensitivity than for dark surfaces is found. The anthropogenic emission sources investigated in this study are located near the Earth's surface, therefore the main effect of clouds is to shield the resulting enhanced SO<sub>2</sub> concentrations from view and thus lower the sensitivity. Therefore, in this study clouds are assumed at an altitude of 4 km. In Table 2 and Fig. 1 (left panel) different parameters are summarized for TRACY-II simulation runs to calculate SO<sub>2</sub> AMF for specific cases; the resulting AMF are presented in Fig. 1 (right panel).

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In order to account for partial cloud cover, the independent pixel approximation (Eq. 1) is used: The SO<sub>2</sub> AMF,  $A$  for the total ground pixel is the average of the SO<sub>2</sub> AMF for its clear and for the cloudy parts (weighted by their geometrical fractions and their respective top of atmosphere radiances).

$$A_{\text{total}} = \frac{A_{\text{cloudy}} \cdot f \cdot I_{\text{cloudy}} + A_{\text{clear}} \cdot (1 - f) \cdot I_{\text{clear}}}{f \cdot I_{\text{cloudy}} + (1 - f) \cdot I_{\text{clear}}} \quad (1)$$

Where  $f$  denotes the effective cloud fraction of the satellite observation, and  $I_{\text{cloudy}}$  and  $I_{\text{clear}}$  the modelled radiances for the cloudy and clear part of the pixel, respectively.

For average effective cloud fraction is assumed 30%, for the observations over Peru and in summertime over Norilsk a surface albedo of 5% is assumed, for the wintertime observations over Norilsk a surface albedo of 80% is assumed. The resulting AMFs for these assumptions are termed as AMF\_CCorrec\_summer and AMF\_CCorrec\_winter respectively, and are also shown as a function of SZA in right panel of Fig. (1). These AMFs are used for the conversion of SO<sub>2</sub> SCD into SO<sub>2</sub> VCD in this study.

It should be noted that for many individual cases, these assumptions might differ from reality. Nevertheless, they might well describe the average situation. Since in this study we are interested in average distributions and relative trends, the detailed choice of the assumed parameters will not critically influence the results. From the comparison of the summer and winter results over Norilsk, Siberia, it is found that the chosen parameters lead to consistent results (see below). Also the good agreement of our data with emissions estimated using other methods confirms the suitability of the chosen parameters (see Sect. 4.5). The resulting SO<sub>2</sub> VCD over the Cu smelters of Peru and Russia with possible influence from other sources and their temporal analysis is presented in the following sections.

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## 4 Results

Significant anthropogenic sources of sulphur emissions, and to a lesser extent nitrogen emissions, exist within the arctic region (AMAP, 2006). Sulfuric acid is the most important acidifying substance in the Arctic, with nitric acid being of secondary importance.

5 Industry situated on the Kola Peninsula and Taimyer Peninsula (Russia), nickel and copper processing works in Nikel, Norilsk, Zapolyarnij and Monchegorsk, are the main contributors to the environmental problems due to the emanating air pollution.

Norilsk and Nikel expanded their presence in international markets during the 1990s, remaining Russia's largest source of air pollution and one of the largest trans-boundary emitters of sulfur in Europe (Barrett, 2000). According to AMAP (1998, 2006), Norilsk, contributes significantly to  $SO_x$  ( $SO_2$ ,  $SO_4^{-2}$ ) emissions above Arctic Circle.

The Cu smelters in Peru are among the most polluting non-volcanic sources of  $SO_2$  in the world. From GOME observations especially the La Oroya and Ilo Cu smelters are found to be strong sources of anthropogenic  $SO_2$  in Peru; they can be clearly identified from 7-years mean map of GOME observations presented in Fig. (4).

### 4.1 1st case study – Norilsk, Siberia, Russia

Since 1971, nickel and copper from the Norilsk ores in Siberia has been processed in smelters. The Norilsk ore contains more sulfur than the ore produced by the Nikel mine (Tommervik et al., 1995; Michelutti et al., 2001; Aamlid, 2002). In total, Norilsk produces over 90 percent of Russia's nickel, 58 percent of copper, over 80 percent of cobalt and almost 100 percent of the platinum-group metals. By the end of 1997, Norilsk exported 95 percent of its copper and nickel (3% of copper and 23% of nickel of the world's total) to the world markets.

Norilsk, Russia's largest exporter of copper, nickel and other non-ferrous metals, is a well-documented example of a major polluter (Kotov and Nikitina, 1996; Michelutti et al., 2001). According to (Carn et al., 2004),  $SO_2$  emissions from Norilsk are variously reported as being on the order of 2 to 3 megatons (Mt) each year (e.g., Foreign

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and Commonwealth Office, Environmental problems in the Russian federation, London Oct 2000, available at: [www.fco.gov.uk/Files/kfile/russiaenviro.pdf](http://www.fco.gov.uk/Files/kfile/russiaenviro.pdf)). However, the emissions estimated in this study from GOME observations over the Norilsk region are  $1.685 \pm 0.3 \text{ Mt SO}_2$  per year (4.61 kt/day on average) are quite close to the range reported by Foreign and Commonwealth Office, Environmental report about Russia (see also Sect. 4.5).  $\text{SO}_2$  emissions from Norilsk smelters ( $69^\circ \text{ N}$ ,  $88^\circ \text{ E}$ ), have been consistently detected by GOME instrument since 1996, as can be clearly identified from 7-years mean map of  $\text{SO}_2$  VCD for 1996–2002 presented in Fig. 2. GOME observations demonstrate that maximum  $\text{SO}_2$  column amounts are found in the vicinity of the Norilsk smelters with a  $\text{SO}_2$  plume extending over the surrounding areas as well.

#### 4.2 Temporal analysis of the $\text{SO}_2$ levels in the Norilsk area

GOME observations over the Norilsk region were analyzed in further detail, and a time series over 7-years is presented in Fig. (3). An area ( $85^\circ \text{ E}$ – $92^\circ \text{ E}$  and  $68^\circ \text{ N}$ – $72^\circ \text{ N}$ ) was selected and the weekly  $\text{SO}_2$  VCD occurring within this region for the whole time period (1996–2002) are presented in Fig. 3. After the Polar night (in Polar Regions, the portion of the year when the Sun does not rise above the horizon. Its length changes from twenty hours at the Arctic/Antarctic Circle (latitude  $66^\circ 33' \text{ N}$  or  $\text{S}$ ) to 179 days at the North/South Pole) between March and October, when the solar zenith angle is  $< 80^\circ$  (as data used in this study are for  $\text{SZA} < 80^\circ$ ),  $\text{SO}_2$  data can be retrieved from GOME observations over Norilsk.

To account for the strong effect of snow on the measurement sensitivity, data about snow cover extension with in selected region ( $85^\circ \text{ E}$ – $92^\circ \text{ E}$  and  $68^\circ \text{ N}$ – $72^\circ \text{ N}$ ) was obtained from Rutgers University Global Snow Lab on special request. Further details about snow coverage data can be obtained from their website (<http://climate.rutgers.edu/snowcover>) or see Robinson and Frei (2000). Transition between snow covered and snow free surface occurs in May, as identified from the snow data. Thus for the trend analysis we skipped the May observations and calculated two separate time series, one for March and April (using the AMF\_CCorrec\_winter) and one for June to

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September (using the AMF\_CCorrec\_summer).

The SO<sub>2</sub> VCD during 1997–1998 are of the same order of magnitude for both, the March–April and June–September data. This indicates that, although the spread of the data points represents the large influence of changing cloud cover, our average assumptions on cloud cover, aerosol properties and surface albedo might be reasonable. Linear fits are applied to both data sets used in Fig. 3 in order to obtain the temporal trend in the SO<sub>2</sub> column amounts. Results demonstrate a slight temporal decrease of 8.7% and 7.3% in GOME SO<sub>2</sub> VCD over Norilsk region for summer and winter months respectively. Similarly AMAP (2006) reported that significant reductions in emissions from the non-ferrous metal smelters on the Kola Peninsula, and to a lesser extent from the Norilsk smelters in the Russian Arctic have been observed over the time period of 1990–2003. Nevertheless, non-ferrous metal productions still remain the dominant source of SO<sub>2</sub> emissions in addition to energy production plants and mining industries in the Arctic region.

#### 4.3 Second case study – Peruvian Cu smelters

Analysis of GOME data revealed two persistent sources of SO<sub>2</sub> in Peru identified as copper smelters situated at La Oroya and Ilo. The La Oroya smelter (11.53° S, 75.90° W) is located in the central western part of Peru and is run by Doe Run Co. The Ilo smelter (17.63° S, 71.33° W) is located at the southern west coast of Peru and is run by Southern Peru Copper Corp (SPCC). Annual capacities of these smelters in 2002 were reportedly 80 and 300 thousand metric tons, respectively (Feliciano and González, 2003), making Ilo one of the ten largest copper smelters in the world. Therefore, relatively strong SO<sub>2</sub> emissions can be observed from the Ilo site in the GOME mean map for the period 1996 to 2002 (Fig. 4).

During the time period of 1996–2002, Reventador and Tungurahua volcanoes from Ecuador were active and resulting SO<sub>2</sub> emissions are clearly visible in Fig. 4. Although, sensitivity of the satellite instruments towards weak boundary layer SO<sub>2</sub> signal (as compared to volcanic SO<sub>2</sub> signal – different plume heights) is usually limited (Eisinger

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and Burrows, 1998; Richter et al., 2002b; Khokhar et al., 2005), our results show that GOME instrument is sensitive enough to be able to identify the SO<sub>2</sub> being emitted from the La Oroya and Ilo smelters. GOME's ability to detect SO<sub>2</sub> emissions efficiently from Peruvian Cu smelters makes it distinct from the other conventional satellite instruments of last decade.

#### 4.4 Temporal analysis over Peru

A temporal analysis of GOME data over the Peruvian Cu smelters, La Oroya and Ilo was made by selecting SO<sub>2</sub> VCD occurring with in these regions: La Oroya (11°–13° S and 75°–78° W) and Ilo (16.5°–18° S and 70°–74° W) for the time period of 1996–2002.

The time series presented in Fig. 5 are averaged over ten days and plotted against the time for whole time period. Linear fits, gray lines, are applied to both data sets in order to obtain the trend in SO<sub>2</sub> emissions. Figure 5a showed a temporal increase of 23% for La Oroya copper smelter, while time series presented in Fig. 5b for Ilo copper smelter showed a temporal decrease of 25% in SO<sub>2</sub> emissions.

The temporal decrease in SO<sub>2</sub> emissions from Ilo smelters might be due to proper emission control and the installation of sulphuric acid plants at the Ilo smelters. According to Boon et al. (2001), in 1997 Ilo's local authorities began implementing the Ilo Clean Air Project, with support from international organizations including the World Health Organization (WHO) and the International Council for Local Environmental Initiatives (ICLEI). A second plant was installed in 1998. These two sulphuric acid plants now convert 30 percent of SO<sub>2</sub> into sulfuric acid (Boon et al., 2001) almost consistent with the observed decrease in GOME SO<sub>2</sub> column amounts. There is no information about installation of such a plant at La Oroya smelter.

#### 4.5 Estimation of the total SO<sub>2</sub> emissions

From satellite observations, total emissions can be estimated by

(a) integrating the SO<sub>2</sub> VCD over an area around the source, and

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(b) assuming an average atmospheric lifetime.

For the first step care should be taken that the emission source of interest is the only significant source within the selected area. The second step, however, is more critical, since the atmospheric lifetime depends on various parameters (e.g. temperature, precipitation, wind) which are difficult to estimate, and which typically vary with time. Nevertheless, in this section we will present a rough estimate of the emissions from the smelters at Norilsk and in Peru. Note that these estimates are completely independent from existing estimates and provide, therefore complementary information. Also from the comparison to existing estimates it is possible to test the assumptions made in Sect. 3. The annual emissions are calculated using the following equation:

$$E_{\text{annual}} = \frac{\int_{\text{area}} V da}{\tau} \cdot 365 \quad (2)$$

Where  $E_{\text{annual}}$  denotes the annual  $\text{SO}_2$  emission,  $V$  the  $\text{SO}_2$  VCD integrated over the selected area, and  $\tau$  the atmospheric lifetime of  $\text{SO}_2$  in days.

We used average VCD for Russian and Peruvian smelters and integrated over area in  $\text{m}^2$  for Norilsk region ( $85^\circ$ – $92^\circ$  E and  $68^\circ$ – $72^\circ$  N), for Ilo region ( $16.5^\circ$ – $18^\circ$  S and  $70^\circ$ – $74^\circ$  W) and for La Orya region ( $11^\circ$ – $13^\circ$  S and  $75^\circ$ – $78^\circ$  W). The distance between given latitude and longitudes coordinates for these regions was calculated by using a tool from Northern Arizona University USA, available online from their website (<http://jan.ucc.nau.edu/~cvm/latlongdist.html>).

The atmospheric lifetime of  $\text{SO}_2$  is highly variable (Stevenson et al., 2003) depending, in particular on altitude of the  $\text{SO}_2$  plume and the presence of clouds (Graf et al., 1997; Khokhar, 2006). It is variously reported ranging from 0.6 to 2.4 days for boundary layer anthropogenic emissions (e.g. Atkinson et al., 1997; Feichter et al., 1997; Graf et al., 1997; Brasseur et al., 1999; IPCC, 2001; Von Glasow et al., 2002; Stevenson et al., 2003; etc.), as compared to volcanic  $\text{SO}_2$  lifetime of 6 days which is injected at high altitudes (e.g. Feichter et al., 1997; Graf et al., 1997; etc.). In this study, lifetime of 1 day for anthropogenic  $\text{SO}_2$  is assumed. After converting  $\text{SO}_2$  VCD (molecules/ $\text{cm}^2$ ) integrated

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over given area i.e. molecules/area into mass/area ( $6.02 \cdot 10^{23}$  molecules equals 64.07 g of  $\text{SO}_2$ ), yearly emissions for the three smelters are derived. For Norilsk smelters,  $\text{SO}_2$  emissions of  $1.685 \pm 0.3$  Mt/yr are comparable to existing estimates 2–3 Mt/yr by Carn et al. (2004). The minor differences could be due to either the assumed lifetime being possibly too long and/or the applied  $\text{SO}_2$  AMF are slightly too large. Similarly,  $\text{SO}_2$  emissions for la Oroya smelter  $0.953 \pm 0.07$  Mt/yr and for Ilo  $1.0953 \pm 0.2$  Mt/yr are close to Carn et al. (2007) estimates for both Peruvian smelters for September 2004–June 2005. However, in general, we conclude that our assumptions made in Sect. 3 are reasonable.

#### 4.6 Regional volcanic eruptions

The South American plate has three principal zones of active volcanism: the Northern Zone (in Ecuador and Columbia), the Central Zone (southern Peru, northern Chile, south-western Bolivia, and north-western Argentina), and the Southern Zone (southern Chile and southern Argentina) (de Silva and Francis, 1991). The northern and southern zones are both relatively well known, at least the identities and locations of the major volcanoes have been documented. By contrast, the Central Volcanic Zone of the Andes remains one of the largest but least known areas of active volcanism in the world. The Central Volcanic Zone of the Andes is located between latitudes  $14^\circ$  and  $28^\circ$  S of the Andean cordillera. This high altitude plateau is mostly consisting of both active and passive volcanoes, much of them over 4000 m in altitude, constituting the Altiplano of Bolivia and Puna of north Chile and Argentina, see Table 3. Since this work focuses on the  $\text{SO}_2$  emissions from Peruvian copper smelters, only the information about Peruvian volcanoes from central Andes volcanic zone are presented. Table 3 gives information about the location, type and the last known activities of Peruvian volcanoes. These information were retrieved from bulletins of global volcanism network (BGVN, <http://www.volcano.si.edu>) and allow investigating any possible influence of volcanic  $\text{SO}_2$  emissions to the Peruvian copper smelter  $\text{SO}_2$  emissions (as it

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happened in case of Hekla eruption – for details see Khokhar, 2006). According to Table 4, only two volcanoes Ubinas and Sabancaya were active during GOME period 1996–2002. Volcanic activities of both volcanoes are discussed in further detail in the following section.

#### 5 4.6.1 Ubinas volcano

The Ubinas volcano (a stratovolcano with 1300 m altitude) is located about 800 km southeast of the La Oroya copper smelter and about 180 km northeast of the Ilo copper smelter. According to Bulletins of GVN (BGVN, 1996), the Ubinas volcano experienced only a minor eruption/degassing during July 1996 and amount of SO<sub>2</sub> was not sufficient to influence the anthropogenic SO<sub>2</sub> from both smelters.

#### 4.6.2 Sabancaya volcano

The Sabancaya volcano (a stratovolcano with 5967 m altitude), is located just 210 km North-West of the Ubinas volcano, about 675 km South-East of the La Oroya and 213 km North-West of Ilo copper smelters.

15 According to bulletins of GVN summarized in Table 3, Sabancaya volcano was active 5 times during the GOME period (1996–2002). It is most probable that any strong volcanic eruption may affect the SO<sub>2</sub> VCD observed over the Ilo copper smelter. Efforts are made to avoid the confusion of volcanic plumes with anthropogenic ones by selecting only SO<sub>2</sub> VCD from area strictly confronting the smelter's SO<sub>2</sub> plume and excluding the volcano location. Furthermore, it is investigated by looking on each of its volcanic activity and extents of volcanic plume precisely and skipping the data for days with volcanic activities of Sabancaya. As stated in BGVN (1997), Sabancaya volcano experienced a moderate eruption in the month of May 1997. On 1st and 2nd of May, aviation reports described ash-bearing plumes. The plume on 1 May reportedly  
20 reached 5.5-km altitude; the one on 2 May, 7.3-km altitude. GOME observed the Sabancaya volcanic eruption on 2nd of May (missed on 1st of May due to a data gap)  
25

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and most likely the data were affected because both, volcano and Ilo copper smelter, typically lie in the same ground pixel (GOME pixel size 320 by 40 km<sup>2</sup>). Therefore, this observation was eliminated from the time series. During July 1997, Sabancaya experienced only fumarolic activity (BGVN, 1997).

5 The GOME instrument observed Sabancaya's volcanic activities efficiently in accordance to GVN bulletins (with the exception of data gaps). However, there was no distinctive evidence found that volcanic SO<sub>2</sub> plume mixed with anthropogenic SO<sub>2</sub> on the basis of the following arguments:

1. Most of the time eruptions were minor and periodic, and lasted only for a few  
10 hours.
2. A maximum plume height of only 1 km was reached except on 2 May 1997. (Data was excluded from analysis)
3. The extent of the volcanic plume was not large enough to reach the south westerly extending plume of Ilo smelter, see Fig. 4. Because this is the region of westerly  
15 prevailing trade winds, so most of the times the plume drifted northwest or north making it less probable that it affected the Ilo plume (situated south east of the volcano).
4. The time series presented in Fig. 5b did not show any correlation of relatively larger peaks with to these volcanic activities in 1997, 1998 and 2000.

20 The only reason why volcanic emissions may still affect the results is the coarse spatial resolution of the GOME. But this can only happen if both volcano and smelter lie in the same ground pixel (taking into account also the effects of finite lifetime and wind transport).

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## 5 Conclusions

Satellite remote sensing systems provide a regular and consistent monitoring of the earth, which is invaluable to monitor the earth system and effect of human activities on it. It is a useful tool to obtain information about various atmospheric trace gas emissions and their distribution on a global scale. The GOME instrument on-board ERS-2 enabled the observations of atmospheric SO<sub>2</sub> for more than 7 years. The analysis of GOME data illustrates that it holds plenty of information on the different SO<sub>2</sub> sources, both of anthropogenic and natural origin.

Anthropogenic SO<sub>2</sub> is much more difficult to detect from satellite instruments than volcanic SO<sub>2</sub> because it is generally located at lower altitudes where the instrument's sensitivity is low (Eisinger and Burrows, 1998; Khokhar, 2006). Enhanced anthropogenic SO<sub>2</sub> emissions can be identified from regions with extensive burning of coal, smelting of metal ores and heavy industrial activities

GOME data showed that the emissions from Norilsk can be clearly identified with maximum SO<sub>2</sub> column amonuts during winter and spring. These high values are mainly caused by underlying snow which enhances the earth's albedo (referred to as albedo effect, see Khokhar et al., 2004, 2006). In this study, efforts are made to take the albedo effect into account by using different parameters for SO<sub>2</sub> AMF calculation during winter and summer. The resulting time series of GOME SO<sub>2</sub> VCD from Norilsk showed a consistent seasonal pattern over the time period of 1996–2002 with an overall slightly decreasing trend. This decrease is in good qualitative agreement with AMAP (2006) and other studies. Nevertheless, the Norilsk smelters are still the dominant source of anthropogenic SO<sub>2</sub> emissions in the region.

The total annual emissions derived from our analysis for different regions are close to that of existing estimates (e.g. Common Wealth, AMAP, 2006; Carn et al., 2004, 2007) indicating that our assumptions on the atmospheric radiative transfer and the SO<sub>2</sub> lifetime are reasonable.

GOME's ability to detect weak signals of boundary layer SO<sub>2</sub> from La Oroya and Ilo

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copper smelters in Peru allowed an efficient monitoring of SO<sub>2</sub> on a regular basis (with exception of data gaps). Time series calculated for La Oroya and Ilo smelters showed a temporal increase by 23% and a 25% decrease in SO<sub>2</sub> emissions from these non-ferrous smelter, respectively over the time period from 1996 to 2002.

5 Due to coarse spatial resolution of GOME, it the SO<sub>2</sub> VCD obtained over Peru could in principle have been affected to some extent by the emissions of the nearby Volcanoes Ubinas and Sabancaya as it happened in case of Norilsk during Hekla eruption in 2000 (for details see Khokhar, 2006). However, from careful sensitivity studies, this potential source of interference could be clearly ruled out. Additionally, GOME obser-  
10 vations are in good agreement with bulletins of global volcanism network (BGVN).

Satellite remote sensing systems with periodic (global coverage in few days) and consistent monitoring have several advantages over other techniques. They also have some limitations, especially coarse spatial resolution (GOME: 40 by 320 km<sup>2</sup>). How-  
15 ever, this limitation will be abated by other satellite instruments onboard space mis- sions from ESA and NASA e.g. SCIAMACHY, OMI or GOME-2, which are already in space and have comparably better spatial resolution. The increased spatial resolution will help to monitor, localize, and study SO<sub>2</sub> plumes more precisely and quantify their impact on a finer spatial scale.

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20 viding of ERS-2, GOME data by DLR (Wessling, Germany) and ESA (Frascati, Italy). We especially acknowledge the GVN Volcanic Emission database from their website <http://www.volcano.si.edu/reports/bulletin/> and for snow coverage data Mr. Thomas Estilow from Rutgers University Global Snow Lab. In addition, special thanks to all colleagues for their contribution and tremendous help. We are particularly grateful to T. Deutschmann who  
25 developed the Monte-Carlo radiative transfer model TRACY-II.

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**Table 1.** Database of Cu smelters from different countries based on their production capacity\*.

Name	Country	Latitude (° N)	Longitude (° E)	Capacity (Ktons/yr)	Process Type
Pirdop	Bulgaria	42.7	24.18	190	Outokumpu Flash
Noranda, Horne, Quebec	Canada	48.25	-79.02	200	Noranda Continuous
Timmins, Ontario	Canada	48.47	-81.33	155	Mitsubishi Continuous
La Negra (Altonorte)	Chile	-23.8	-70.33	165	Noranda Continuous
Chuquicamata	Chile	-22.32	-68.93	500	Outokumpu/Teniente Conv.
El Teniente (Caletones)	Chile	-34.1	-70.33	400	Reverberatory/Teniente Conv.
Daye, Hubei	China	30.08	114.95	150	Noranda Conv./Reverberatory
Guixi, Jiangxi	China	28.28	117.18	250	Outokumpu Flash
Kunming, Yunnan	China	25.04	102.72	230	Isasmelt, Electric
Lunen (Huettenwerke Kayser)	Germany	51.62	7.52	170	Issmelt
Hamburg (Norddeutsche Affinerie)	Germany	53.55	10	420	Outokumpu, Contimelt, Electric
Besshi/Toyo, Ehime	Japan	33.88	133.32	300	Outokumpu Flash
Naoshima, Kagawa	Japan	34.45	134	270	Mitsubishi Continuous
Onahama, Fukushima	Japan	36.95	140.9	348	Reverberatory
Saganoseki, Oita	Japan	33.24	131.87	470	Outokumpu Flash
Tamano, Okayama	Japan	34.48	133.93	263	Outokumpu Flash
Nacozari de Garcia, Sonora (La Caridad)	Mexico	30.37	-109.53	300	Outokumpu/Teniente Conv.
La Oroya	Peru	-11.53	-75.9	80	Reverberatory
Ilo	Peru	-17.63	-71.33	300	Reverberatory/Teniente Conv.
Glogow District (Glogow I)	Poland	51.67	16.08	220	Blast Furnace
Glogow District (Glogow II)	Poland	51.67	16.08	205	Outokumpu Flash
Kirovgrad	Russia	57.44	60.06	150	Reverberatory (S)
Norlisk	Russia	69.34	88.22	400	Reverb and Electric
Phalaborwa	South Africa	-23.98	31.07	135	Reverberatory
Huelva	Spain	37.26	-6.95	320	Outokumpu Flash
Skelleftehamn (Ronnskar)	Sweden	65	21.6	240	Electric (TBRC)
Garfield, Utah	USA	37.75	-111.5	320	Kennecott/Outokumpu
Hayden, Arizona	USA	33	-110.78	210	Inco Flash
Miami, Arizona	USA	33.38	-110.87	180	Isasmelt/Electric

(RLE)=Roast Leach, (S)=secondary smelter, (TBRC)=Top Blown Rotary Converter

(\*) table adapted from website <http://pubs.usgs.gov/of/2003/of03-075/CSTable.txt>

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**Table 2.** TRACY-II scenarios of SO<sub>2</sub> AMF calculations.

SO <sub>2</sub> AMF	Wavelength (nm)	Aerosol profiles	SSA	Asymmetry parameter “g”	Surface Albedo %
noAerosol	315	–	–	–	5
Aerosol	315	see Fig. 1	0.95	0.68	5
Clouds 80%	315	Clouds at 4 km	1	0.85	80
Clouds 50%	315	Clouds at 4 km	1	0.85	50
Clouds 5%	315	Clouds at 4 km	1	0.85	5

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**Table 3.** Database of Peruvian Volcanoes – summarized from bulletins of GVN.

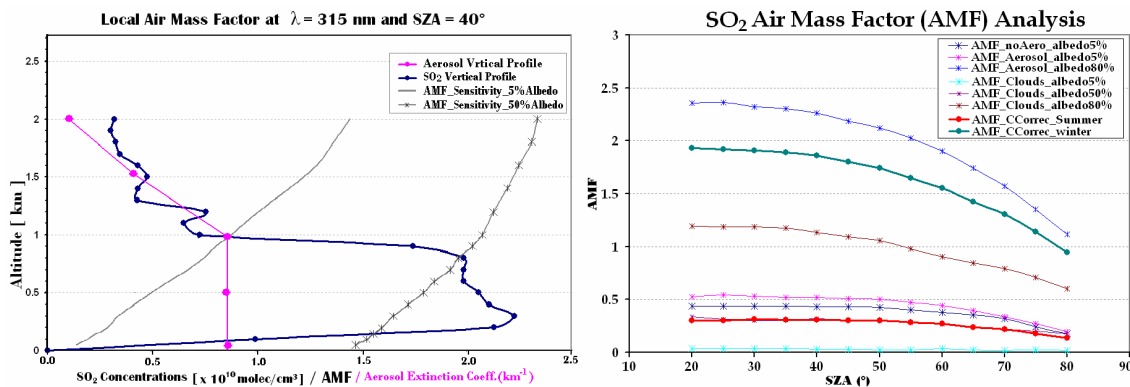
Volcano	Location	Elevation (m)	Volcano type	Last Known Eruptions*
ANDAHUA VALLEY	15.42° S–72.33° W	4713	Cinder cones	unknown
AUQUIHUATO, CERRO	15.07° S–73.18° W	4980	Cinder cone	unknown
CASIRI, NEVADOS	17.47° S–69.81° W	5650	Stratovolcano	unknown
CHACHANI, NEVADO	16.19° S–71.53° W	6057	Stratovolcano	unknown
COROPUNA	15.52° S–72.65° W	6377	Stratovolcano	unknown
HUAYNAPUTINA	16.608° S–70.85° W	4850	Stratovolcano	1600
EL MISTI	16.29° S–71.40° W	5822	Stratovolcano	05/1984, 12/1985
QUIMSACHATA	14.20° S–71.33° W	3923	Lava dome	4450 BC (?)
SABANCAYA	15.78° S–71.85° W	5967	Stratovolcano	06/1988, 05/1990, 06/1990, 07/1990, 05/1991, 07/1991, 01/1992, 03/1994, 05/1995, <b>07/1997, 05/1998 08/1998 10/1998</b> , 05/2000, 05/2003, 01/2004
SARA SARA	5.33° S–73.45° W	5522	Stratovolcano	unknown
TICSANI	16.75° S–70.59° W	5408	Lava domes	unknown
TUTUPACA	17.02° S–70.35° W	5815	Stratovolcano	unknown
UBINAS	16.35° S–70.90° W	5672	Stratovolcano	06/1969, 07/1969, 12/1985, 07/1996, 03/2006, 05/2006, 10/2006, 2007
YUCAMANE	17.18° S–70.20° W	5550	Stratovolcano	1902

(\*): Information retrieved from GVN Bulletins, see references

(?): not certain

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**Fig. 1.** Left Panel: Vertical profiles of the SO<sub>2</sub> concentration and aerosol extinction coefficient used for AMF calculation. Also shown is the measurement sensitivity (expressed as Box-AMF) towards different surfaces with albedos of 5% (solid Gray line) and 50% (Gray line with cross). Right panel: comparison of SO<sub>2</sub> AMFs for different scenarios. Also shown are selected SO<sub>2</sub> AMF used in this study: AMF\_CCcorrec\_summer (red thick line) used for Norilsk in summer and for Peru, AMF\_CCcorrec\_winter (greenish thick line) used for Norilsk in winter.

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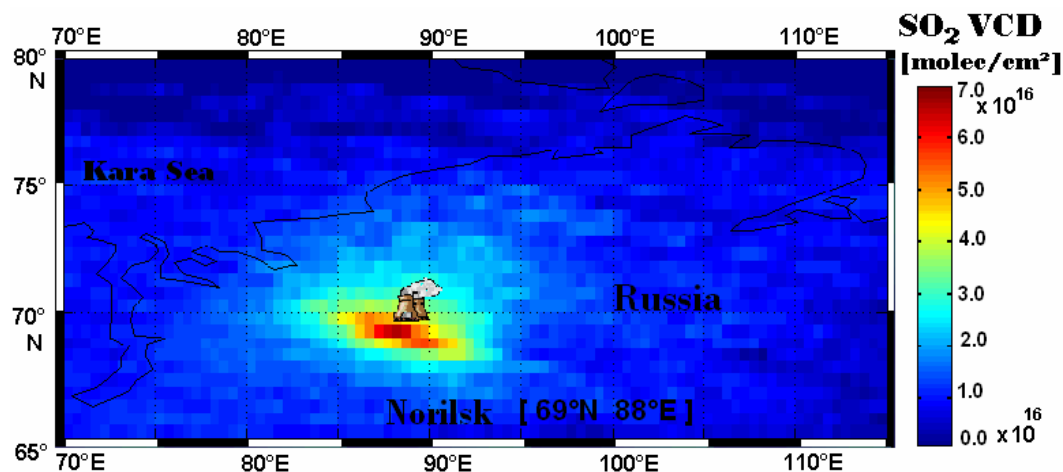
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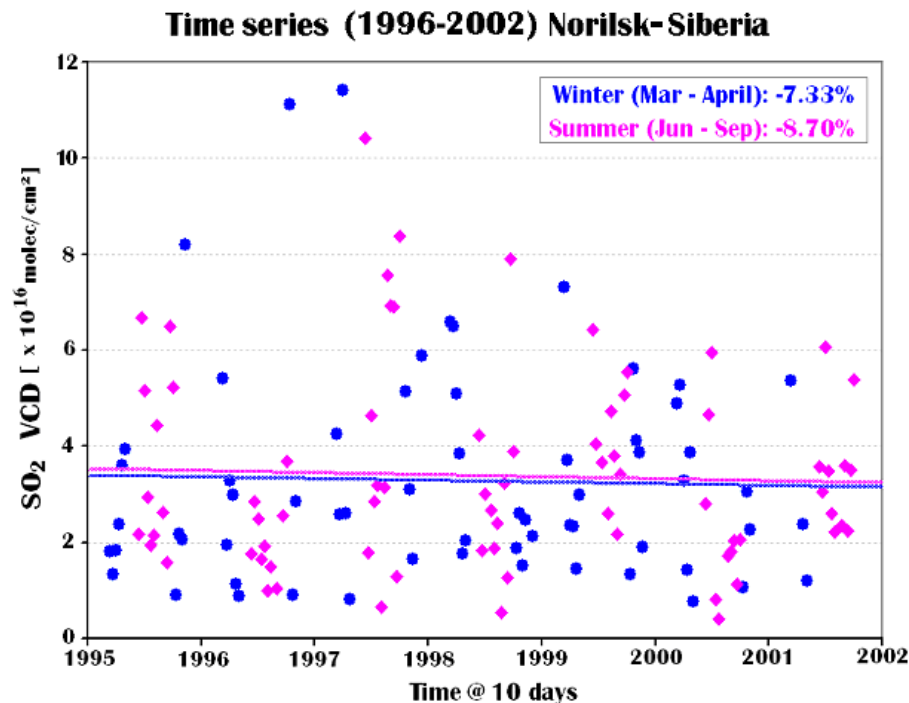


**Fig. 2.** 7-year (1996–2002) mean map of the SO<sub>2</sub> vertical column density over Norilsk, Russia, showing a strong source of anthropogenic SO<sub>2</sub> emissions. The location of Norilsk copper smelters is indicated by the smoke-stack symbol. Two different AMFs are used for winter and summer months to prepare this mean map (see text).

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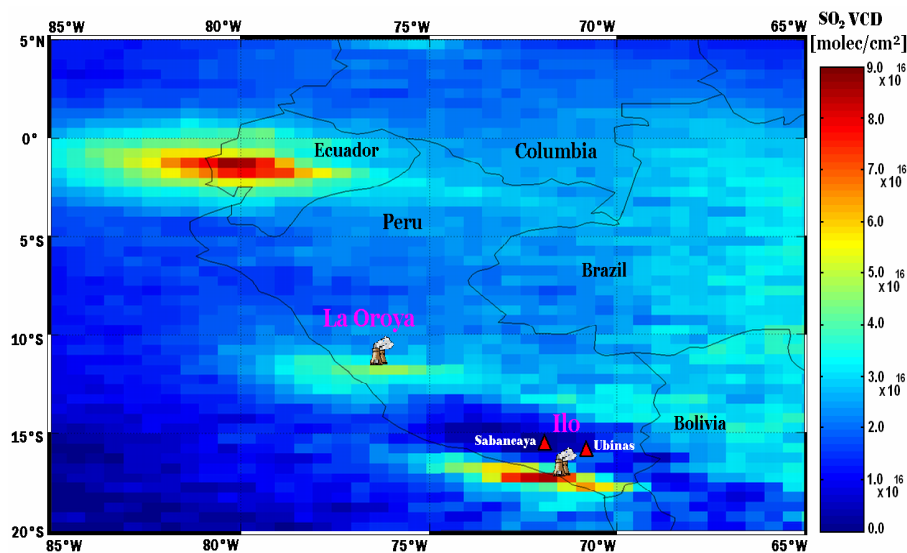


**Fig. 3.** Time series of the area averaged SO<sub>2</sub> VCD retrieved from GOME data (1996–2002) over Norilsk. Linear fits are applied in order to identify the temporal trend. The spread of the data is mainly caused by variations in cloud cover.

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**Fig. 4.** 7-years (1996–2002) mean map of SO<sub>2</sub> VCD retrieved from GOME data (1996–2002) over Peru and surrounding region. In addition to the hotspots caused by the La Oroya and Ilo Peruvian Cu smelters (indicated by smoke-stack symbols), enhanced SO<sub>2</sub> VCD can also be clearly identified over Ecuador (due to volcanic eruption during this time period).

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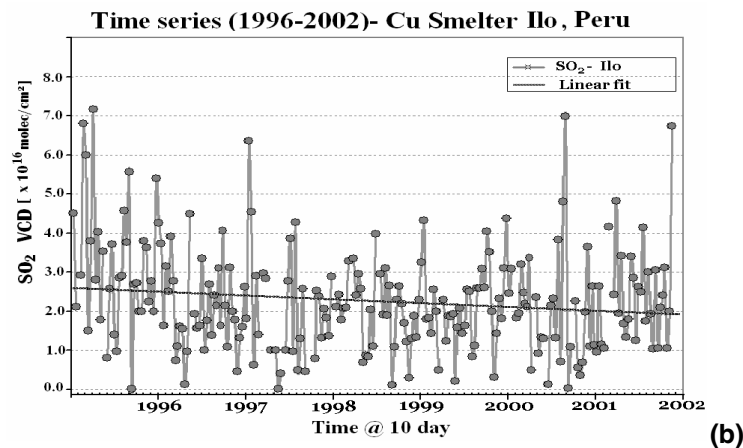
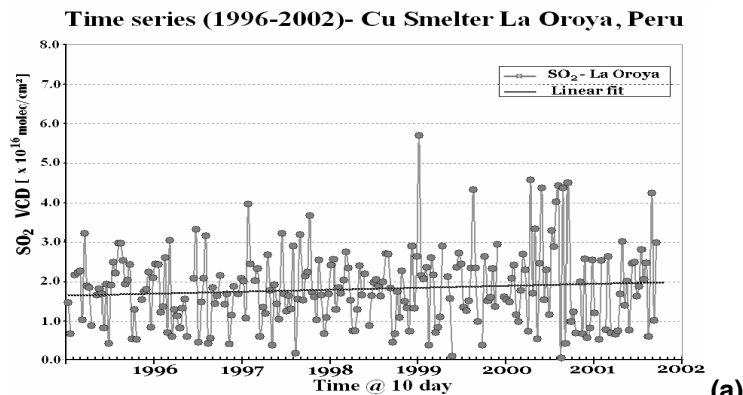
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**Fig. 5.** Time series of the area averaged SO<sub>2</sub> VCD over selected regions over La Oroya (a) and Ilo (b) Cu smelters, retrieved from GOME data from 1996 to 2002. The individual data points are averages over ten days. Linear fits (Black lines) are applied to both time series. Time Series over La Oroya Cu smelters (a) indicates an increase of 23% while time series over Ilo Cu smelters (b) indicates a decrease of 25%.

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