

Evaluation of the MERIS aerosol product over land with AERONET

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Abstract. The Medium Resolution Imaging Spectrometer (MERIS) launched in February 2002 on-board the ENVISAT spacecraft is making global observations of top-of-atmosphere (TOA) radiances. Aerosol optical properties are retrieved over land using Look-Up Table (LUT) based algorithm and surface reflectances in the blue and the red spectral regions. We compared instantaneous aerosol optical thicknesses retrieved by MERIS in the blue and the red at locations containing sites within the Aerosol Robotic Network (AERONET). Between 2002 and 2005, a set of 500 MERIS images were used in this study. The result shows that, over land, MERIS aerosol optical thicknesses are well retrieved in the blue and poorly retrieved in the red, leading to an underestimation of the Angstrom coefficient. Correlations are improved by applying a simple criterion to avoid scenes probably contaminated by thin clouds. To investigate the weakness of the MERIS algorithm, ground-based radiometer measurements have been used in order to retrieve new aerosol models, based on their Inherent Optical Properties (IOP). These new aerosol models slightly improve the correlation, but the main problem of the MERIS aerosol product over land can be attributed to the surface reflectance model in the red.

1 Introduction

There is a clearly need for an accurate representation of the distribution of aerosols over the globe not only because of their direct and indirect radiative impacts on climate (IPCC, 2007), but also because of their health impact on population (Wilson and Sprengler, 1996). The representation of the aerosols optical properties distribution is provided by several tools, from satellite aerosols products (Kaufman et

al., 2002), surface measurements (Dubovik et al., 2002) and aerosol transport model (Chin et al., 2002). Information about aerosol absorption is often needed for radiative impact purposes, but it is still difficult to accurately obtain this quantity at global scales with current space sensors (Mishchenko et al., 2004). From space, actual retrievals on aerosol optical properties are mainly based on three different techniques: (i) from multi-bands unpolarized measurements, (ii) with polarization and/or (iii) multidirectionality. All of these different techniques provide advantages/inconveniences on the aerosol retrieval. For example, multi-bands unpolarized sensors like the Moderate Resolution Imaging Spectroradiometer (MODIS) sensor allow a good spatial resolution at the ground but can provide information on total column amount and size of aerosols (Remer et al., 2005). The aerosol parameters have been recently improved with the MODIS Second Generation Algorithm (Levy et al., 2007) and the “Deep-Blue” algorithm (Hsu et al., 2006). These new algorithms enhanced the possibility to discriminate dust particles from fine aerosols. Using the multidirectionality as the Multiangle Imaging Spectroradiometer (MISR) sensor, provide constraints on the surface reflectance and on scattering properties of aerosols (Abdou et al., 2005). Adding the polarized measurements like POLDER increases the information content and provides constraints on the surface reflectance and on the fine mode of the aerosol distribution (Deuzé et al., 2001).

The Medium Resolution Imaging Spectrometer (MERIS) instrument can also assume an integral role in the effort of obtaining a global picture of aerosols due to its frequent global measurements of aerosol amount and type over a wide variety of surface types. The primary goal of MERIS is the ocean color observation, while the secondary purpose is the observation of the atmosphere and the terrestrial surface. MERIS is one of the instruments of the ENVISAT satellite launched in 2002. ENVISAT is a sun-synchronous orbit with an equator crossing time of 10:00 a.m. local time. MERIS is a programmable, medium-spectral resolution,



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Table 1. Biome number, model name, associated AERONET site names and Principal Investigator (PI) with range of τ_a at 440 nm and Angstrom coefficient α over the number of match-ups N .

Biome	Model name	AERONET sites	PI	$\tau_a(440)$ range	α range	N
0	Boreal America	Bonanza Creek	J. Hollingsworth	0.06–1.96	0.6–1.9	33
		Bratts Lake	B. McArthur	0.05–0.58	1.1–2.0	29
		Pickle Lake	B. McArthur	0.04–0.21	1.3–1.8	5
1	Boreal Euroasia	Andenes	B. Holben	0.07–0.31	1.3–1.9	3
		Yakutsk	M. Panchenko	0.07–0.26	0.6–2.2	8
2	MidLat West America	Rimrock	B. Holben	0.14–0.79	1.7–2.0	4
3	MidLat East America	GSFC	B. Holben	0.03–1.23	0.9–2.3	59
		Bondville	B. Holben	0.06–1.41	0.4–2.1	32
		Walker Branch	B. Holben	0.06–0.70	1.0–2.3	23
4	MidLat Europe	Minsk	A. Chaikovsky	0.08–1.47	0.9–1.9	23
		Lille	P. Goloub	0.07–0.89	0.3–1.8	37
		Ispira	G. Zibordi	0.02–1.12	0.6–3.4	87
5	MidLat Asia	Beijing	H.-B. Chen	0.13–2.96	0.7–1.5	29
		Tomsk	M. Panchenko	0.09–0.39	1.6–2.0	12
6	Tropical America	Tenosique	M. J. Montero-Martinez	0.1–2.22	0.9–2.0	19
7	Tropical Asia	Bac Giang	H. Vet Le	0.34–1.4	0.9–1.6	11
		Pimai	B. Holben	0.35	0.3	1
8	Equatorial America	Alta Floresta	B. Holben	0.05–1.26	0.4–2.1	48
		Campo Grande Sonda	E. B. Pereira	0.05–0.31	0.7–2.0	6
		Belterra	B. Holben	0.08–0.34	0.7–1.4	8
9	Equatorial Africa	Mongu	B. Holben	0.04–0.31	0.9–2.7	18
		Ilorin	R. T. Pinker	0.59	0.7	1
		Djougou	P. Goloub	0.28–0.5	0.4–1.2	3
10	Equatorial Asia	Jabiru	R. Mitchell	0.08–0.24	0.2–1.0	6

imaging spectrometer operating in the solar reflective spectral range. Fifteen spectral bands can be selected by ground command, each of which has a programmable width and a programmable location in the 390 nm to 1040 nm spectral range. The instrument's 68.5° field of view around nadir covers a swath width of 1150 km with a spatial resolution of 1.2 km at nadir (Rast et al., 1999). The MERIS accuracy is $\pm 4\%$ in reflectance (Delwart et al., 2003). The absolute uncertainties of the vicarious calibration of MERIS over land are found between 3 and 7%, depending on the accuracies of the available ground truth data (Kneubuehler et al., 2004).

In this article, we evaluate the MERIS aerosol product over land. The first part will be devoted to the presentation of the general aspects of the aerosol retrieval over land from multi-channel sensors working in visible (VIS) to near infrared (NIR) spectral regions. Both 1st and 2nd MERIS processing are described. The second part presents the worldwide ground-based aerosol measurement Aerosol Robotic Network (AERONET) sites that we used in the evaluation

of the MERIS aerosol product over land. The third section will compare the MERIS aerosol product over land against AERONET outputs, then describes a new aerosol models family based on AERONET sky radiances measurements. The new aerosol model family obtains Inherent Optical Properties of aerosols (IOPA) that slightly improved the MERIS aerosol product over land. Lastly, we will point out the weakness of the surface reflectance that explains the poor MERIS aerosol product over land in the red.

2 MERIS aerosol retrieval

2.1 Generality of the MERIS aerosol remote sensing over land

Aerosol remote sensing over land from space is a very difficult task because of the high reflectivity of the Earth compared to the aerosol scattering signal in the back-scattering region. The technique chosen for MERIS (Santer et al.,

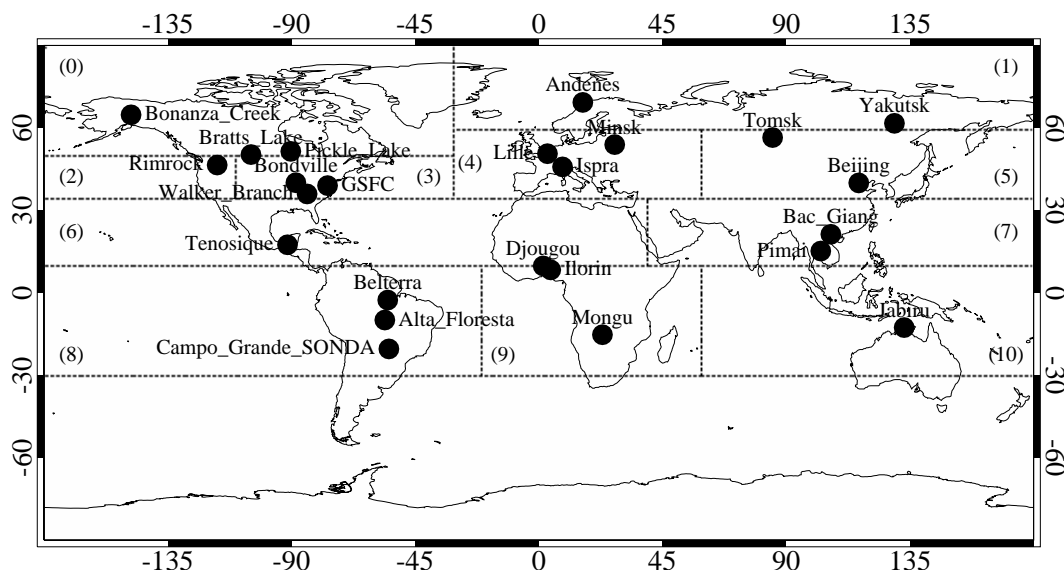


Fig. 1. Geographical distribution of the 24 AERONET sites used in this study. Biomes are represented by dashed boxes with associated numbers given between brackets.

1999) relies on the well known Dense Dark Vegetation (DDV, Kaufman and Sendra, 1988) concept, generalized to the dark target concept of MODIS (Kaufman et al., 1997). The idea here is to detect dark and stable targets whose reflectivity is known accurately with a simple and reliable method. For MERIS, the Atmospherically Resistant Vegetation Index (ARVI, Kaufman and Tanré, 1992) is used to detect DDV. ARVI is defined by

$$\text{ARVI} = \frac{\rho_{ag}^{865} - (\rho_{ag}^{670} - \gamma(\rho_{ag}^{443} - \rho_{ag}^{670}))}{\rho_{ag}^{865} + (\rho_{ag}^{670} - \gamma(\rho_{ag}^{443} - \rho_{ag}^{670}))}, \quad (1)$$

where ρ_{ag} is the reflectance at 865 nm, 670 nm or 443 nm, corrected from gaseous absorption and Rayleigh scattering. γ is a coefficient fixed to 1.3 for DDV (Santer et al., 1999). ARVI allows obtaining a relation between NIR band (at 865 nm where aerosols contribution is low) and red and blue bands (at 670 nm and 443 nm respectively, where surface contributions are low). By comparison, the MODIS team uses the capability to observe in the near infrared at $2.1 \mu\text{m}$ for detecting dark targets. In all of these techniques, the two main sources of uncertainties are the accuracy of the reflectance model of the target and the accuracy of the aerosol models, which were used for the computation of the aerosol scattering functions.

2.2 The aerosol retrieval in the 1st MERIS processing

In the case of MERIS, 11 biomes have been chosen to represent the spatial and the temporal variations of the DDV concept over the globe. Figure 1 gives the geographical distribution of the 11 biomes (represented by dashed boxes with the number in brackets, from 0 to 10). Table 1 provides the

model name of each biome. For each biome a set of Look Up Tables (LUT) has been generated that gives the DDV Bidirectional Reflectance Function (BRDF) in the three bands (blue, red and NIR) and the coupling terms between the DDV and the atmosphere (Ramon and Santer, 2001). The aerosol characterization is based on aerosol models. They are defined by a size distribution $n(R)$ of particles of radius R represented by the Junge power law, $n(R) \approx R^{-(\alpha+3)}$, and by 26 values of the Angstrom coefficient α_j (from 0 to 2.5 by step of 0.1). These models will be called Junge models hereafter. They are also defined by 3 values for the real part of the refractive index m (1.33, 1.44 and 1.55). No absorption has been included in the aerosol models. At the present time, the aerosol refractive index is set to 1.44 by default, which corresponds to a standard continental aerosol model. The aerosol optical properties (extinction coefficient, single scattering albedo and phase function) have been precalculated at 550 nm with the Mie theory. Aerosol optical thicknesses (AOTs) τ_a in the red and in the blue are retrieved for each Angstrom coefficient α_j . The model for which the Angstrom coefficient is the closest to the one that is obtained from the τ_a retrieval is the model that we select. The aerosol product of the 1st MERIS processing consists of τ_a at 865 nm and α .

2.3 The aerosol retrieval in the 2nd MERIS processing

Because the concept of DDV was too restrictive in order to get a good spatial representation of the aerosols, the model of the surface reflectance has been extended to brighter surfaces called Land Aerosol Remote Sensing (LARS). The LARS surface reflectances in the blue and in the red that have a linear dependence with the ARVI (Santer et al., 2007a), can

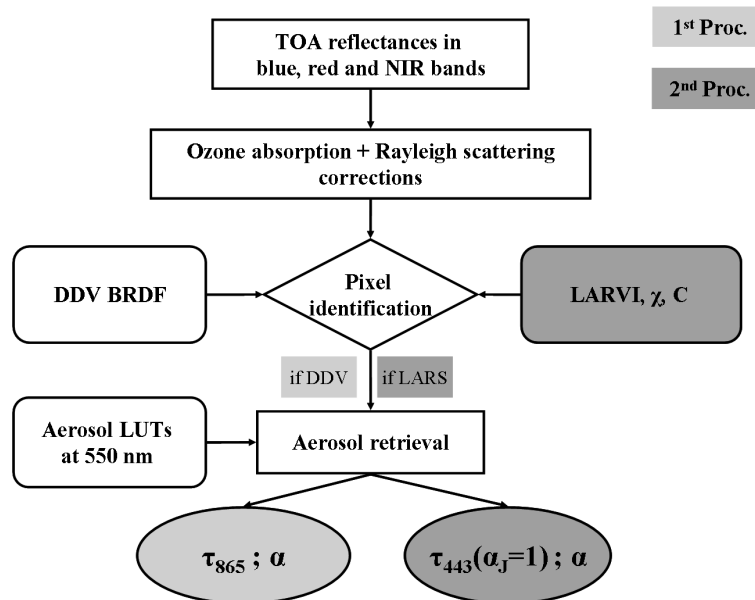


Fig. 2. Flowchart of the MERIS aerosol retrieval over land. Grey boxes show the change between the 1st processing and the 2nd processing.

be written as

$$\rho_{\text{LARS}}(m, \lambda, \theta_v, \theta_s, \phi_v - \phi_s) = C(m) \rho_{\text{DDV}}(\lambda, \theta_v, \theta_s, \phi_v - \phi_s) (1 + \chi(\lambda, m)(\text{LARVI} - \text{ARVI})), \quad (2)$$

and depends on the month (m), the wavelength (λ), the viewing zenith angle (θ_v), the solar zenith angle (θ_s), and the azimuth angle difference between viewing and solar ($\Phi_v - \Phi_s$). Further details of the estimation of the monthly coefficients (C), the slopes of the linear dependency χ and the lower values of the ARVI (LARVI) where the linear dependence is valid, can be found in Santer et al. (2007a).

The aerosol product from the 2nd MERIS processing offers a much better spatial coverage but also introduces more errors in the τ_a and α retrievals. These errors occur mostly in the red, as the variation of the surface reflectance with the ARVI is more pronounced than in the blue and therefore most subject to uncertainties. Preliminary tests of the aerosol retrieval using the LARS indicated a large and random spatial distribution of α rendering suspect the retrieval of τ_a in the red. Efforts were made to improve the characterization of the surface reflectance in the blue and in the red using the MODIS level 3 albedo maps (Moody et al., 2003) to produce the required surface reflectance. Both the offset and slope of the linear dependence with the ARVI, in a 1° by 1° spatial grid, have been defined (Ramon and Santer, 2005). The initial 11 biomes are kept in order to describe the BRDF of the LARS pixels. Nevertheless, α values were still suspicious. They are flagged because out of range values [0–2.5] on numerous occasions. Therefore, the following decisions were taken: (i) to output τ_a at 443 nm instead of τ_a at 865 nm because of the disastrous effect of α ; (ii) to produce the τ_a in the blue using the Junge model of $\alpha_J=1$ and (iii) to output α as an indicator

of the aerosol type as it was computed for the 1st processing. Then, the aerosol product of the 2nd MERIS processing consists of τ_a in the blue (at 443 nm) and α . Figure 2 provides the flowchart of both processings.

3 AERONET data

AERONET is a globally distributed network of automated ground-based instruments and data archive system, developed to support the aerosol community. The instruments used are CIMEL spectral radiometers that measure the spectral extinction of the direct Sun radiance (Holben et al., 1998). The aerosol optical depths are determined using the Beer–Bouguer Law in several spectral bands. For this study, level-2 data are used and consist of τ_a at 440 nm and 675 nm, retrieved at least every 15 min during day time. Level-2 data are cloud-free and quality assured retrieved from pre- and post-field calibrated measurements (Smirnov et al., 2000). The estimated accuracy in the AERONET τ_a is between ± 0.01 and ± 0.02 and depends on the wavelength, for an air-mass equal to 1 (Dubovik et al., 2000).

We selected geographically diverse AERONET sites that provided generally good-quality measurements records between 2002 and 2005. A total of 24 AERONET sites have been selected in order to cover the 11 biomes of MERIS and the range of possible aerosol optical thicknesses from clean areas to turbid ones (due to different air masses types and sources as biomass burning, continental and/or dusty conditions). Figure 1 gives the geographical distribution of AERONET sites (represented by black dots). We optimally selected three AERONET sites per biome. Unfortunately,

biomes 2 (MidLat West America), biome 6 (Tropical America) and biome 10 (Equatorial Asia) are under represented with only one site because of the lack of AERONET sites, the lack of AERONET measurement or the area not covered by vegetations (mainly desert or snow-covered areas). Information about AERONET sites per biome (name and Principal Investigator of the site) are provided in Table 1. We also gave the range of τ_a in the blue and the range of α over the number of match-ups (N) that have been used in this study.

Biome 0 (Boreal America) is represented by three AERONET sites that cover τ_a in the blue from 0.04 to 1.96 and α between 0.6 and 1.9. The large value of α is representative of small particles. The arctic atmosphere is generally clear but frequently subjected to forest fires in Alaska. Jet streams also transport pollution from Asia or other source regions into this region (Bokoyé et al., 2002). Biome 1 (Boreal Euroasia) is represented by two AERONET sites but suffers from a lack of match-ups ($N=11$) and low variability of τ_a in the blue (from 0.07 to 0.31). Those regions can also be subjected by long range transport of arctic haze (Toledano et al., 2006) or forest fires, that explains the high values of the α (up to 2.2). Biome 2 (Mid Latitude West America) is represented by only one AERONET site and few match-ups ($N=4$) where τ_a in the blue is between 0.14 and 0.79. This part of North America can be affected by aerosols transported from Eurasia. Biome 3 (Mid Latitude West America) is represented by three AERONET sites with τ_a in the blue from 0.03 to 1.41. Biome 4 (Mid Latitude Europe) is represented by three AERONET sites with τ_a in the blue from 0.02 to 1.47. AERONET sites of biomes 3 and 4 are continental sites that cover a diversity of urban and industrial pollution aerosols (Kahn et al., 2005). Biome 5 (Mid Latitude Asia) is represented by two AERONET sites with τ_a in the blue from 0.09 to 2.96. Due to combined influences of arid dust region production and increased fossil fuel usage, the East Asia regions often experience very high concentrations of tropospheric aerosols (Eck et al., 2005). Biome 6 (Tropical America) is represented by one AERONET site with τ_a in the blue from 0.1 to 2.22. The Mexico area is considered as a heavily urban polluted site. Biome 7 (Tropical Asia) is represented by two AERONET sites with τ_a in the blue from 0.34 to 1.4. Those different sites are industrialized urban area (Grey et al., 2006). Biome 8 (Equatorial America) is represented by three AERONET sites with τ_a in the blue from 0.05 to 1.26. The Amazonian Basin is a great source of biomass burning aerosol during the period from August to October (Schafer et al., 2002). Biome 9 (Equatorial Africa) is represented by three AERONET sites with τ_a in the blue from 0.04 to 0.59. Africa is an important source of desert dust and biomass burning aerosols (Eck et al., 2001). Biome 10 (Equatorial Asia) is represented by one AERONET site with τ_a in the blue from 0.08 to 0.24.

The different sites we selected will give us a good picture of the quality of the MERIS aerosol optical depths over land. However, for some biomes, we do not expect to make any

conclusion on the quality of the MERIS aerosol retrieval due to the lack of match-ups (such as biomes 2 and 10) but we kept them for the global comparison nevertheless.

4 The results

4.1 Initial validation

In order to take into account both the spatial and temporal variability of aerosol distribution, the MERIS level-2 aerosol product at 1.2 km pixel resolution and the AERONET direct Sun measurements need to be collocated in space and time. We required at least 2 out of possible 5 AERONET measurements within ± 30 min of MERIS overpasses and at least 10% out of possible MERIS retrievals in a square box of 10×10 pixels centered over AERONET sites (that represent 10 measurements over a 12×12 km² area). The mean values of the collocated spatial and temporal ensemble are then used in linear regression and root mean square errors (rmse) analysis. The total number of match-ups we obtained was 500 for the 24 AERONET sites between 2002 and 2005. The left panel of Table 2 give an overview of the results between 2nd processing MERIS and AERONET τ_a in the blue. The number of match-ups N , the correlation coefficient r , the linear regression equation coefficients (slope and intercept) and the rmse are provided for each biome. For biomes 1 and 10, correlations are poor (with correlation coefficient of 0.23 and 0.37, respectively) certainly due to a wrong surface reflectance model in these extreme areas. The correlation for biome 2 is perfect ($r=1$) but biased by the few match-ups. For others biomes, correlations are good with at best $r=0.93$ for biome 0. In most cases, slopes are greater than 1, which implies an underestimation of the MERIS τ_a compared to the AERONET value. But in some cases, we are very close to the 1:1 line (for example, see biomes 0 and 4). Intercepts are small and rmse are comprises between 0.139 and 0.53. The latter high value of rmse of 0.53 for biome 5 might be explained by an effect of absorption that is not taken into account in our aerosol models. One particular feature that we can observe in some cases is that MERIS shows a very large value of τ_a when compared to AERONET. This might be explained by the presence of thin clouds, like cirrus, that the actual MERIS algorithm is not able to flag. In order to remove these contaminated scenes, we applied a simple threshold on the standard deviation of τ_a in the blue within the box (called σ -filter hereafter). A value of 0.15 seems to be the best value (D. Ramon, personal communication) in order to remove inhomogeneous scene contaminated by thin clouds. We applied the σ -filter and in the right panel of Table 2, we provided the statistical outputs from the σ -filtered match-ups scatterplots. In most biomes, correlation coefficients slightly increased, rmse decreased without significantly changing neither the slope nor offset coefficients of the linear regression. However, the σ -filter is not the most

Table 2. Results of the comparison between the 2nd processing MERIS τ_a and AERONET τ_a in the blue for each biome for the initial match-ups (unfiltered) and for the σ -filtered match-ups.

Biome	N	r	Unfiltered			σ -filtered				
			Slope	Intercept	rmse	N	r	Slope	Intercept	rmse
0	67	0.93	1.03	-0.09	0.149	60	0.97	1.10	-0.09	0.110
1	11	0.23	0.22	0.07	0.204	3	0.98	1.34	0.18	0.088
2	4	1.00	1.86	-0.39	0.139	3	0.97	2.08	-0.46	0.135
3	114	0.89	1.34	-0.11	0.147	104	0.90	1.38	-0.11	0.141
4	142	0.77	1.05	-0.02	0.142	130	0.82	1.11	-0.03	0.127
5	41	0.62	0.82	-0.09	0.530	32	0.61	0.85	-0.12	0.550
6	19	0.98	1.41	-0.18	0.187	19	0.98	1.41	0.18	0.187
7	12	0.83	1.36	-0.12	0.242	12	0.83	1.36	-0.12	0.242
8	62	0.89	1.11	-0.12	0.184	54	0.93	1.18	-0.12	0.155
9	22	0.59	0.6	0.05	0.135	20	0.70	0.87	-0.01	0.109
10	6	0.37	0.13	0.12	0.250	4	0.98	0.47	0.05	0.120

Table 3. Results of the comparison between the 2nd processing MERIS τ_a and AERONET τ_a in the red for each biome for the initial match-ups (unfiltered) and for the σ -filtered match-ups.

biome	N	r	Unfiltered			σ -filtered				
			Slope	Intercept	rmse	N	r	Slope	Intercept	rmse
0	67	0.90	0.76	-0.06	0.166	60	0.94	0.83	-0.07	0.138
1	11	0.09	0.05	0.06	0.222	3	1.00	0.88	-0.12	0.147
2	4	1.00	0.77	-0.14	0.230	3	0.97	1.50	-0.35	0.206
3	114	0.74	0.86	-0.04	0.127	104	0.76	0.91	-0.04	0.117
4	142	0.60	0.59	0.02	0.134	130	0.64	0.65	0.01	0.120
5	41	0.57	0.48	-0.01	0.532	32	0.55	0.50	-0.03	0.523
6	19	0.94	1.00	-0.11	0.147	19	0.94	1.00	-0.11	0.147
7	12	0.78	1.04	-0.12	0.167	12	0.78	1.04	-0.12	0.167
8	62	0.86	0.60	-0.04	0.227	54	0.89	0.62	-0.03	0.206
9	22	0.32	0.27	0.05	0.182	20	0.46	0.50	0.01	0.145
10	6	0.41	0.14	0.08	0.250	4	0.97	0.44	0.03	0.138

efficient test to remove high thin cirrus because (i) they can be very homogeneous over the scene and (ii), they may have an optical thickness lower or near the same order of magnitude as aerosols'. In order to minimize the effect of high, thin cirrus clouds, oxygen pressure can be utilized as an effective mask. Indeed, MERIS offers the possibility to accurately retrieve the cloud top pressure thanks to its window and absorption dual channels in the O₂ A-band (Preusker and Fischer, 1999). The use of the cloud top pressure retrieval is even able to mask very thin cirrus clouds (Ramon et al., 2002). Unfortunately, we did not use the cloud top pressure based mask in our study.

The same comparison has been done in the red. Table 3 provides the summary of statistical outputs from the scatterplots both with and without the σ -filter (i.e., the threshold on the standard deviation of τ_a in the blue). Main conclusions of the comparison are that MERIS overestimates τ_a compared

to AERONET, and the correlations are reduced in most of the cases in the red than in the blue. In the red, the application of the filter allows us to improve correlations.

In order to summarize the comparison, we combine all the data irregardless of their location. Figure 3 indicates the quality of the MERIS retrieval in the blue from the 2nd processing without (Fig. 3a) and with (Fig. 3b) the σ -filter. Correlations in the blue are very good with $r=0.8$ and a linear regression close to the 1:1 line (slopes of 0.98 and small negative intercept of -0.03). The σ -filter allows to slightly increase the correlation coefficient to 0.83 and to reduce the rmse from 0.215 to 0.2. Figure 4 indicates the quality of the MERIS retrieval in the red from the 2nd processing without (Fig. 4a) and with (Fig. 4b) the σ -filter. Overall, the MERIS τ_a retrieval in the red is not as good as the retrieval in the blue, as $r=0.7$ (increased to 0.73 with the σ -filter) and with the slope of the regression of 0.57 (increased to 0.62 with

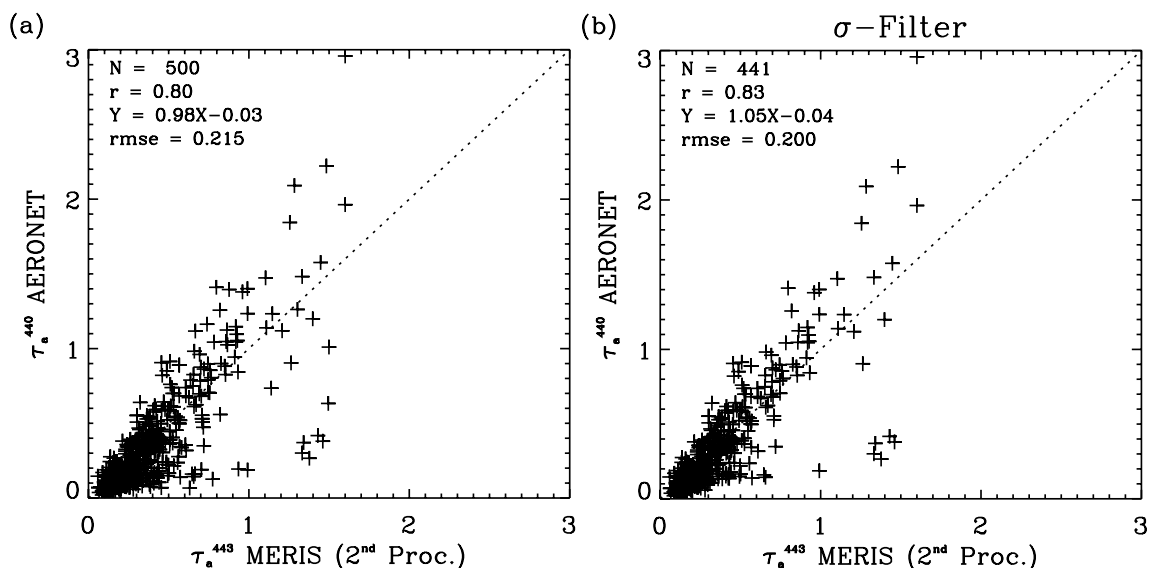


Fig. 3. Scatterplot of τ_a AERONET versus 2nd processing τ_a MERIS in the blue for the initial match-ups (a) and for the σ -filtered match-ups (b).

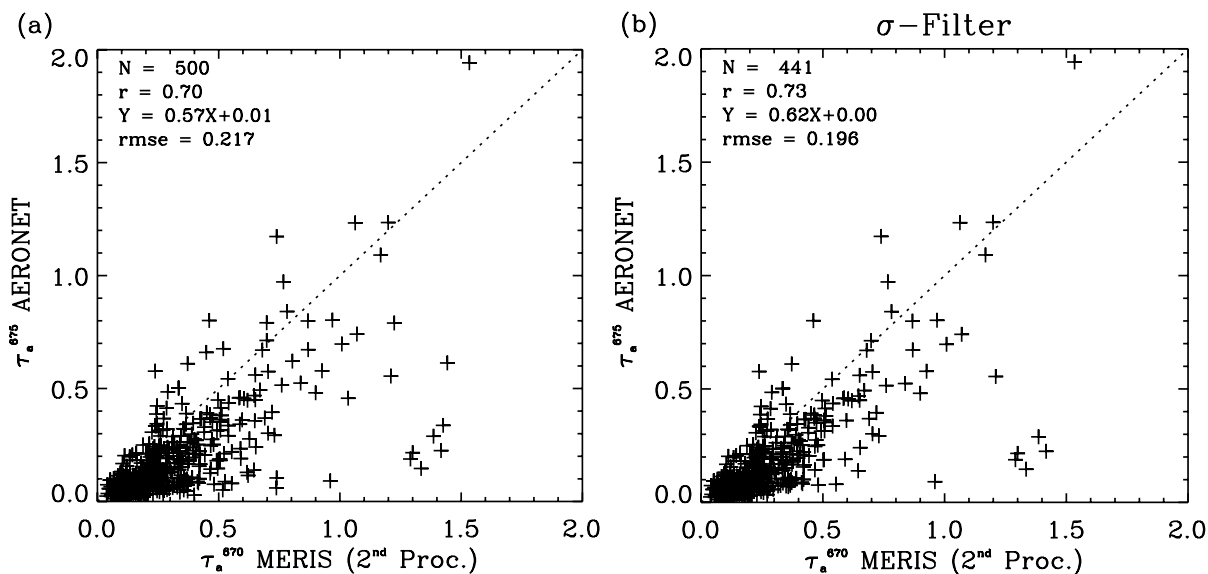


Fig. 4. Same as Fig. 3 for τ_a in the red.

the σ -filter). In Fig. 5, the similar comparison on α is shown and looks less favorable. In both cases (i.e., for match-ups selected with or without the σ -filter), MERIS shows larger aerosols than AERONET.

4.2 Relevance of the 2nd MERIS processing approach

At this stage, we have only presented the evaluation of the MERIS 2nd processing aerosol product over land. We can alter the Junge models in order to evaluate the 1st processing. Actually, in the 2nd processing, it is partially the 1st processing in the combined retrieval of τ_a in the blue and in the red (or τ_a in the blue and α). The main difference is that it is applied to LARS pixels in the 2nd processing, whereas it was applied to DDV pixels in the 1st processing. But α remains unchanged between the two processing on LARS pixels.

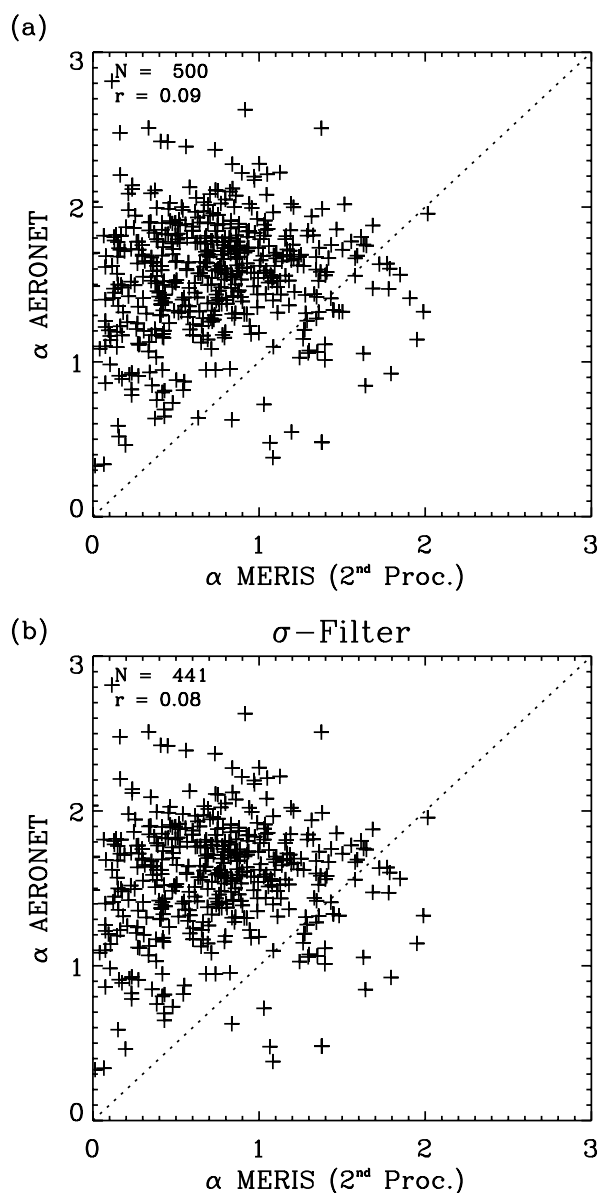


Fig. 5. Same as Fig. 3 for α .

In the 1st processing, α was used as the spectral dependence of τ_a in the retrieval of τ_a in the blue. Then to evaluate the 1st processing, we have to recalculate τ_a from the 2nd processing aerosol product (i.e., τ_a and $\alpha_J=1$) by using the retrieved α . Figure 6 gives the results of τ_a in the blue on the σ -filtered match-ups. Taking the retrieved α compared to $\alpha=1$ leads to a depreciated retrieval of the τ_a in the blue. The correlation coefficient decreased from 0.83 (2nd processing) to 0.72 (1st processing). The slope also decreased from 1.05 to 0.73, respectively. The rmse increased by 0.071.

In order to explain the depreciation, we have to introduce the relation between the aerosol path radiance (L_a) that MERIS measures and the aerosol product τ_a . In the pri-

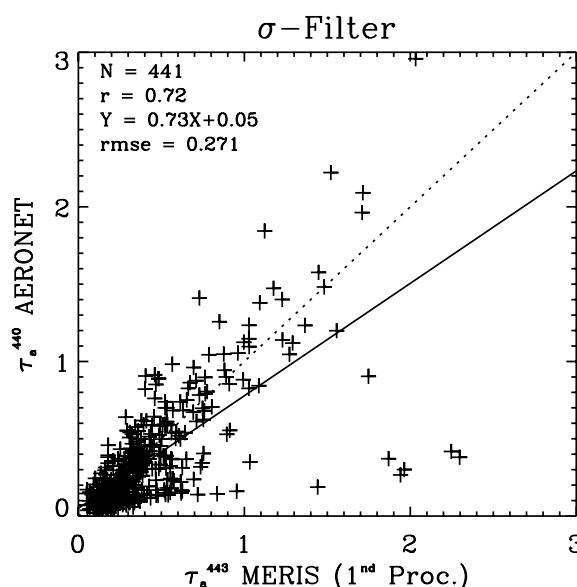


Fig. 6. Scatterplot of τ_a AERONET versus 1st processing τ_a MERIS in the blue for the σ -filtered match-ups.

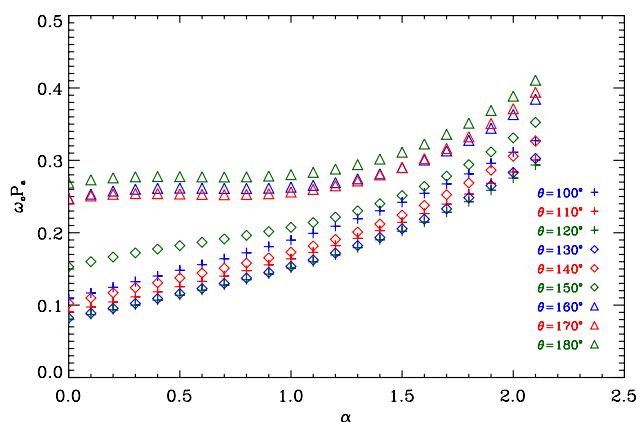


Fig. 7. The dependency of $\omega_0 P_a$ versus α in the MERIS LUTs.

mary scattering approximation, the aerosol path radiance is expressed by:

$$L_a^0 = \frac{\tau_a \omega_0 P_a(\theta) E_s}{4\mu_v \pi}, \quad (3)$$

where E_s is the solar irradiance for the central wavelength of each spectral band corrected for the Sun-Earth distance, ω_0 is the single scattering albedo, P_a the aerosol phase function, θ the scattering angle and μ_v the cosine of the sensor viewing angle.

The depreciated retrievals can now be explained by two effects. First, the product $\omega_0 P_a$ in the backscattering region (for MERIS θ is comprise between 100° to 150°) increases with α . Figure 7 show $\omega_0 P_a$ versus α for different values of θ . Secondly, MERIS underestimates α (see Fig. 5b); therefore MERIS underestimates $\omega_0 P_a$ and overestimates τ_a in

the blue. The decision to arbitrary set $\alpha=1$ for the MERIS 2nd processing is justified.

4.3 Doing better with a new set of aerosol models?

The interpretation of the aerosol path radiance into τ_a relies on the use of 26 Junge models. The ability for these aerosol models to describe the aerosol optical properties was reported by Ramon et al. (2003). This validation exercise was based on CIMEL AERONET measurement of the sky radiance in the principal plane. The following methodology, described by Santer and Lemire (2004), was used:

- α between 440 nm and 675 nm is used to select the two boundary Junge models,
- The Successive Order of Scattering (SOS, Deuzé et al., 1989) code is used to simulate the sky radiance in the principal plane,
- The inputs to the SOS code are the CIMEL τ_a at the time of the sky radiance measurements and with the corresponding geometrical conditions,
- Simulated and measured sky radiances are compared.

This evaluation of the Junge models led to noticeable discrepancies in the sky radiance retrieval. A similar approach highlights that the Junge models overestimated the sky radiance with a systematic bias of 10% at 675 nm and 30% at 870 nm (Aznay and Santer, 2007). The performance of the Junge models in the blue is a bit more difficult to conclude because of the predominance of the Rayleigh scattering.

An alternative use of the CIMEL sky radiance is to retrieve $\omega_0 P_a$ (Santer and Martiny, 2003; Santer et al., 2007b). Therefore, the European Space Agency (EAS) undertook an action to produce a new set of aerosol models based on the interpretation of the CIMEL sky radiances to retrieve the inherent optical properties of aerosols, i.e. the product $\omega_0 P_a$ (called IOPA models hereafter). This new set of aerosol models is still classified in α for values between 0 and 2.5 by step of 0.1. Results are reported in Zagolski et al. (2007) for a similar approach conducted over water. We shown in Fig. 8, the comparison of $\omega_0 P_a$ for different α between the initial values using the Junge models and the IOPA models retrieved from AERONET. The agreement is sometimes excellent, mainly for values of α near 1 (Fig. 8c, d and e).

4.3.1 The MERIS 2nd processing with IOPA models: deriving τ_a in the blue

In the 2nd processing, the $\alpha=1$ Junge model is selected, so we do not expect to see spectacular changes on the retrieval of τ_a in the blue when replacing the Junge aerosol models by the IOPA models. In order to change the aerosol model, we can at first use the primary scattering approximation to describe the aerosol path radiance (Eq. 3). If we change $\omega_0 P_a$, then

we use a simple ratio technique to derive a new value of τ_a , that is given by:

$$\frac{\tau_a^{\text{IOPA}}}{\tau_a^{\text{Junge}}} = \frac{[\varpi_0 P_a(\theta)]_{\text{Junge}}}{[\varpi_0 P_a(\theta)]_{\text{IOPA}}}. \quad (4)$$

But if we modify τ_a , then we have to take into account the change in the multiple scattering factor f , defined as the ratio between primary scattering and multiple scattering, i.e.,

$$L = f L_a^0 = f \frac{\tau_a \varpi_0 P_a(\theta) E_s}{4\mu_v \pi}, \quad (5)$$

In the MERIS ground segment, the multiple scattering factor f has been implemented in the form of LUT computed with the SOS code and generated for the 26 Junge models. To derive τ_a with the IOPA, we need to reconstruct the aerosol path radiance. It can be done through Equation 5 with the former Junge models. But the interpretation of Eq. 5 with IOPA models requires the knowledge of f . It is not so simple to generate a set of new LUTs of f with the IOPA models. We choose, according to some hypothesis, to use f implemented in the MERIS ground segment with the IOPA models. The main hypothesis that we made is that f is the same for two families of aerosol models. To validate this hypothesis, we simulated f with the SOS code for two different families of aerosol models, the Junge models and the Shettle and Fenn (1979) models corresponding to the same α . In Fig. 9, we plotted f at 870 nm for three classes of aerosol models (coastal, maritime and rural) with a relative humidity of 50%. The solar zenith angle was set to 70° and τ_a was set to 0.15. As we can see, there is no big difference between the two families, particularly as to the rural model between 110° and 150° of θ . The comparison of the new τ_a with the IOPA models is reported in Fig. 10. The quality of the linear regression is slightly improved; with a slope of 1.01 for IOPA models and 1.05 with the Junge models (Fig. 3b). These changes are not considered to be significant.

4.3.2 The MERIS 1st processing with IOPA model: deriving α and τ_a in the blue

In this part we explored the possibility to return to the 1st processing with the IOPA models. Starting from the aerosol path radiance in the blue, we used the α derived from MERIS and its associated IOPA models that give the $\omega_0 P_a$ to derive τ_a in the blue. With α , we obtained τ_a in the red. That is the 1st processing and its associated τ_a values. Then, we reconstructed the aerosol path radiance in the red as we did in the blue. We could then vary both aerosol path radiances. So, we used the MERIS algorithm that is described as follows:

1. A double loop is done with 26 α_J (26 Junge models) and with τ_a to retrieve the aerosol path radiance. Outputs are 26 values of τ_a .
2. This double loop is applied in the blue and in the red. The resulting τ_a are reported in Fig. 11.

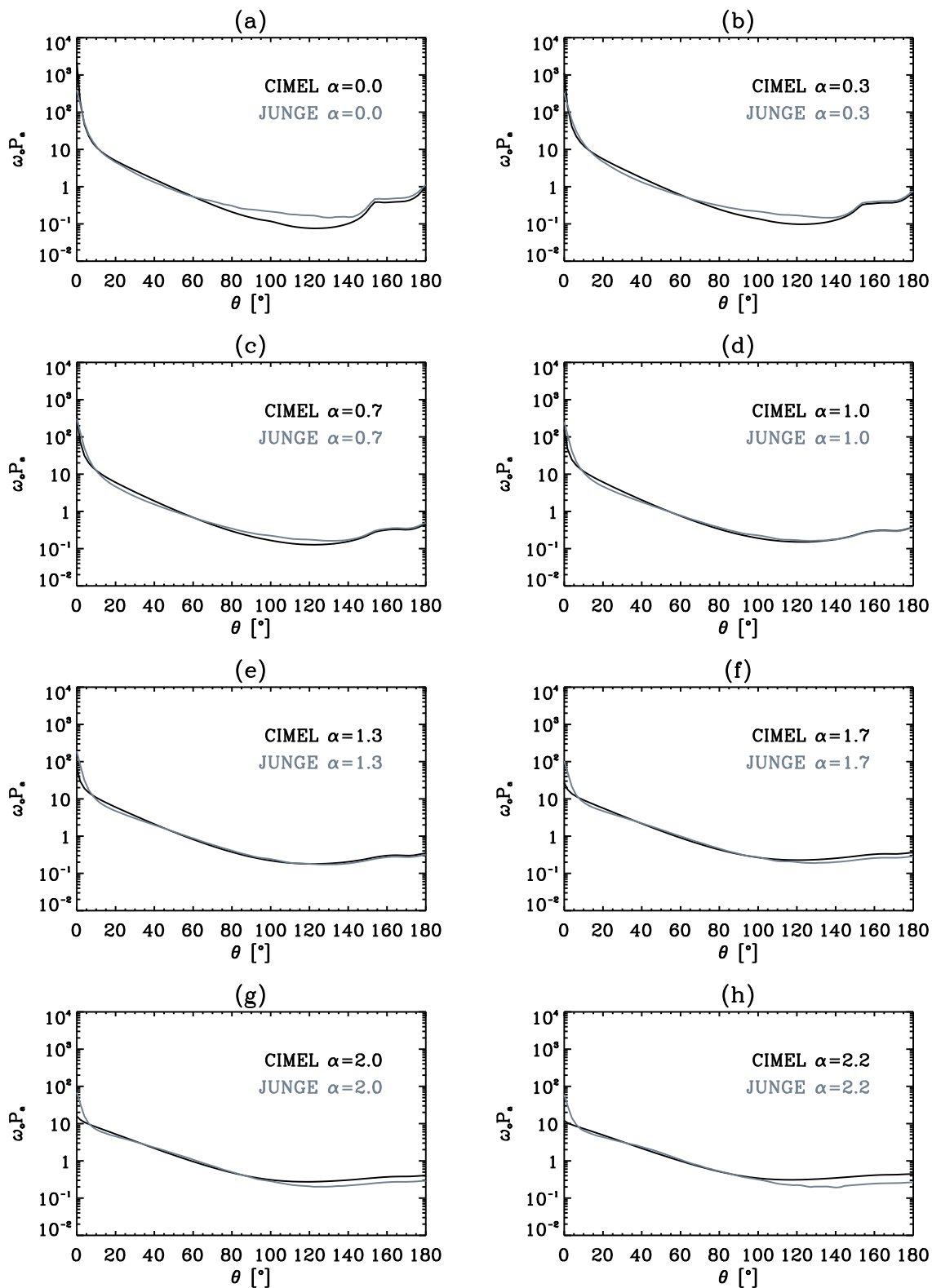


Fig. 8. Comparison on $\omega_0 P_a$ versus the scattering angle θ between the initial values using the Junge models and the retrieved values from CIMEL sky measurements. Results are reported for different values of α (from a to h).

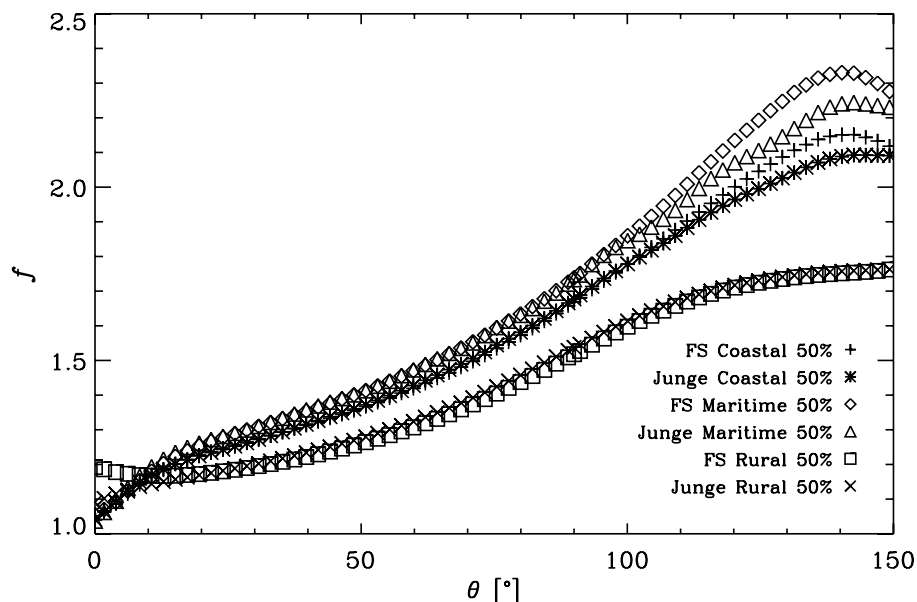


Fig. 9. Multiple scattering factor f at 870 nm versus the scattering angle θ calculated for two different aerosol families (Junge models and Fenn and Shettle (FS) models) and for three aerosol types (coastal, maritime and rural) with a relative humidity of 50%.

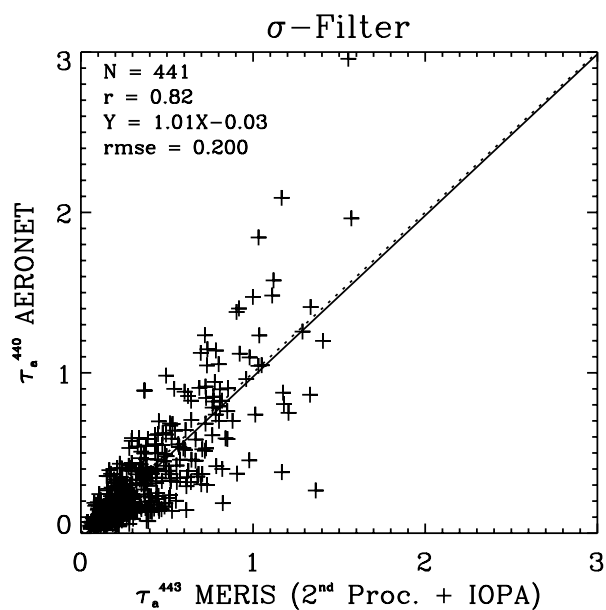


Fig. 10. Scatterplot of τ_a AERONET versus 2nd processing τ_a MERIS retrieved with the IOPA models in the blue for the σ -filtered match-ups.

3. From the two series of retrieved τ_a , we compute an Angstrom coefficient α_{MERIS} .
4. When $\alpha_{\text{MERIS}} = \alpha_J$, we get the final τ_a .

Some comments about Fig. 11 are necessary. When α increases, $\omega_0 P_a$ increases, as we can see on Fig. 7. Therefore,

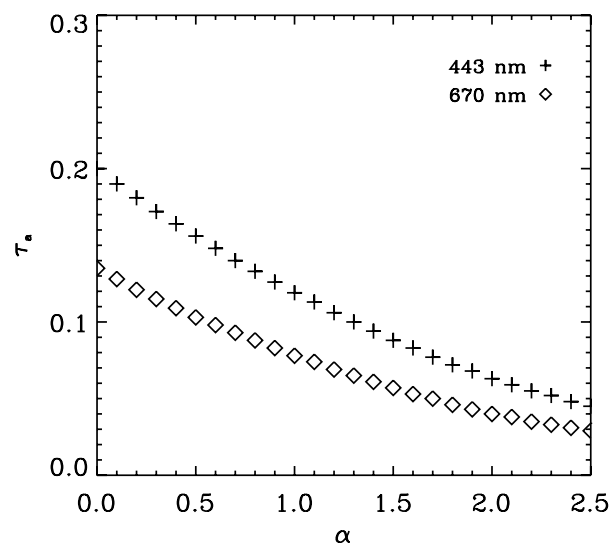


Fig. 11. τ_a in the blue and in the red versus α as obtained by the MERIS 1st processing by looping on the 26 Junge models. These results have been obtained on the Alta Floresta AERONET site on June, 18th 2002.

τ_a decreases with α in proportion with $\omega_0 P_a$ in the primary scattering approximation (Eq. 3), and it does a little more when accounting from the multiple scatterings. Because the aerosol phase function has no wavelength dependency, the τ_a ratio 443/670 is insensitive to α , if we exclude the second order effect of the multiple scattering. We can now apply the MERIS first processing with IOPA models. Figure 12 illustrates the comparison of τ_a in the blue (left panel) and in

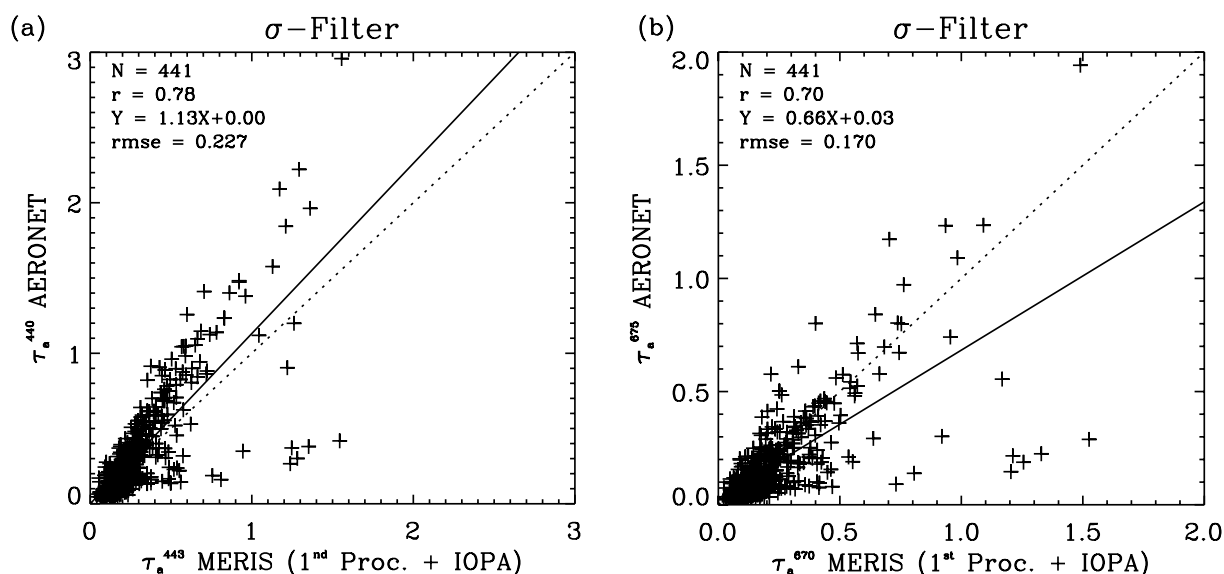


Fig. 12. Scatterplot of τ_a AERONET versus 1st processing τ_a MERIS retrieved with the IOPA models in the blue (a) and in the red (b) for the σ -filtered match-ups.

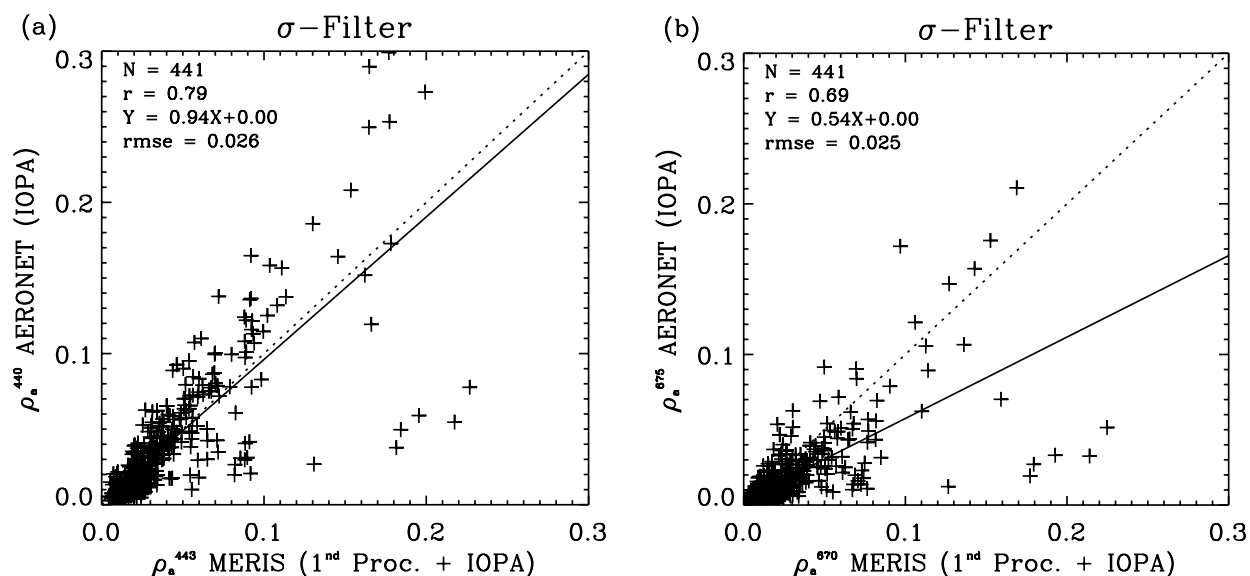


Fig. 13. Scatterplot of AERONET aerosol reflectance versus 1st processing MERIS aerosol reflectance recalculated with the IOPA models in the blue (a) and in the red (b) for the σ -filtered match-ups.

the red (right panel) with the AERONET outputs. Introducing the IOPA models into the MERIS 1st processing leads to a slight increase of the correlation coefficient in the blue from 0.72 (Junge models) to 0.78 (IOPA) and a decreasing of the rmse from 0.271 to 0.227, respectively. However, the 2nd processing with IOPA gives better results at least in the blue (Table 4). In the red, we still have an overestimation of τ_a compared to AERONET.

4.4 Possible errors in the LARS reflectance at 670 nm?

With the aerosol models, the other key parameter in the MERIS τ_a retrieval over land is the knowledge of the LARS surface reflectance. We expect that an inaccuracy in the LARS surface reflectance has less impact on the τ_a retrieval in the blue when compared to the red for the following reasons: (i) vegetation appears darker in the blue than in the red

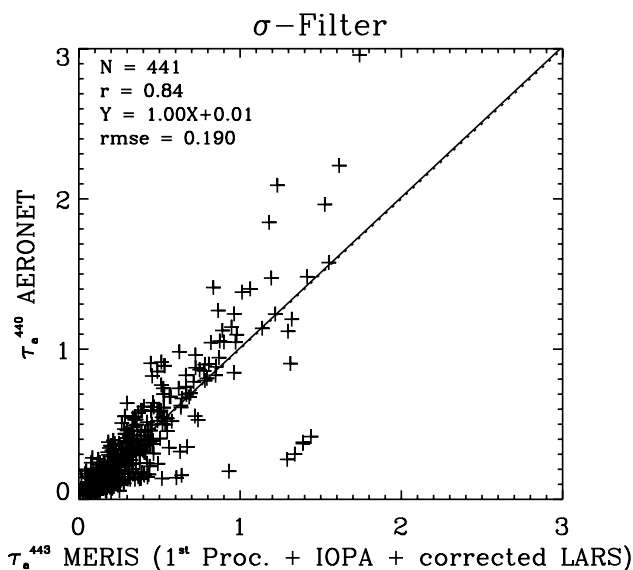
Table 4. Summary of the statistical outputs from the scatterplots applied on the σ -filtered match-ups for the different cases studied here.

case	blue				red			
	r	Slope	Intercept	rmse	r	Slope	Intercept	rmse
1st proc./Junge models	0.72	0.73	0.05	0.271	–	–	–	–
2nd proc./Junge models	0.83	1.05	–0.04	0.200	0.73	0.62	0	0.196
1st proc./IOPA models	0.78	1.13	0	0.227	0.70	0.66	0.03	0.170
2nd proc./IOPA models	0.82	1.01	–0.03	0.200	–	–	–	–
2nd proc./IOPA models + corrected LARS	0.84	1.00	0.01	0.190	–	–	–	–

and (ii) the aerosol signal increases towards the blue. It is clearly the case from the above results (Table 4). To effectively demonstrate this, let us assume that the aerosol type is known. First, let us take α as measured by AERONET. If we have the correct aerosol model with IOPA, then we should have the correct aerosol reflectance if the LARS reflectance is correct. We ran the MERIS 1st processing with the “exact” aerosol type and output the aerosol reflectance. Results are reported in Fig. 13. The retrieval in the blue is a little biased and remains bad in the red, due to the LARS reflectance. In the blue, the MERIS LARS reflectance is a little high, resulting in an under determination of τ_a . Conversely, in the red, the MERIS LARS reflectance is too low resulting in an over determination of τ_a . The two combined give an underestimate of α . Now if we correct the LARS reflectance in the blue based on the underestimation of the aerosol reflectance, we are able to retrieve a new τ_a in the blue. Figure 14 shows the comparison of the τ_a AERONET versus the τ_a MERIS in the blue from the IOPA models and the corrected LARS reflectance. We finally obtained a slight increase of the correlation coefficient to 0.84 with a slope of the linear regression equal to 1 with a very small intercept of 0.01.

5 Conclusion and recommendations

An extensive data set of CIMEL AERONET measurements was used in the evaluation of the MERIS aerosol product over land. This aerosol product consists basically in τ_a in the blue and in the red. There is, at first, a clear need to better filter the MERIS τ_a within the box selected for the comparison between MERIS and AERONET. The filtering used here was based on the spatial homogeneity of τ_a with a threshold on the standard deviation within the box. Artificial spatial variations of τ_a are commonly caused by, (i) the wrong cloud masking: it is known that cirrus clouds are badly detected and the use of the O₂ A-band would be very useful, (ii) the edges of a cloud: Santer et al. (2005) noticed artificial increases of τ_a in the vicinity of clouds, and (iii) the shadow of the cloud: in the MERIS processing, LARS pixels in the cloud shadows are rejected by a radiometric threshold which has to be validated. One solution to overcome these difficulties is to

**Fig. 14.** Scatterplot of τ_a AERONET versus 1st processing τ_a MERIS retrieved with the IOPA models and the corrected LARS surface reflectance in the blue for the σ -filtered match-ups.

supervise the selection of the validation points. This painful process will reduce the number of validation points. Clearly a validation of the aerosol product has to be conducted on a daily level 3 product. This level 3 product should be elaborated taking into account the different origins of the biases in the τ_a retrieval.

After a simple data filtering (based on a threshold of 0.15 on the standard deviation of τ_a in the blue over the 10×10 pixels box), the first conclusion is that MERIS correctly estimates τ_a in the blue compared to AERONET with a regression equation of $y=1.05x-0.04$ and a correlation coefficient of 0.83. The application of this filter is recommended to any given user who wants to use the data because we found that the filter reduced the rmse from 0.215 to 0.2 in the blue and from 0.217 to 0.196 in the red. However, α is clearly strongly underestimated. By spectral extrapolation, we can imagine the disaster for τ_a at 865 nm, which is the standard product over ocean, and then on the homogeneity between water-land products. The reference in τ_a in the blue

is therefore relevant.

The reconstruction of the 1st processing on LARS pixels instead of DDV pixels indicates that it is better to arbitrary set α to 1. It justifies the choice made for the 2nd processing which allows to propose a better spatial coverage of the aerosol product combined to a reliable estimate of τ_a in the blue. Furthermore, we found that the rmse is reduced from 0.271 for the 1st processing to 0.2 for the 2nd processing showing that the current version of the MERIS aerosol product over land is the most appropriate.

The choice of a Junge models to describe the aerosol optical properties was quite arbitrary. It was sustained by the simplification in the LUT generation due to the non dependence of the aerosol phase function with wavelength. It did not pretend to describe the microphysical properties of the aerosol by their inherent optical properties.

Using alternative aerosol models based on the CIMEL sky radiance measurements at AERONET sites, the IOPA models, confirms that the main improvement necessary concerns the LARS surface reflectance in the red. At present time, the LARS surface reflectance LUTs were produced by the MODIS surface albedo map. Alternatively, we can also use the MERIS surface albedo map (Schroeder et al., 2005). It remains that the production of albedo maps requires to apply atmospheric correction, therefore knowing τ_a . This infernal loop is broken if the albedo maps are only produced for clear days, which is difficult to obtain globally.

One alternative to avoid the difficulty in the red is to evaluate the performance of using the couple (412 nm–490 nm) instead of (443 nm–670 nm). A negative result of the aforementioned alternative is the reduced spectral interval. On a positive note, though, is that the LARS reflectance at 490 nm is slightly darker than at 670 nm. Moreover, the linear dependency of the LARS reflectance at 490 nm versus the ARVI is less pronounced than at 670 nm. It is foreseen to use the MERIS prototype to test alternative LUTs of LARS reflectance as well as to combine 412 nm and 490 nm on these AERONET match-ups.

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