

Optical properties and mineralogical composition of different Saharan mineral dust samples: a laboratory study

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Abstract. In aerosol chamber experiments optical properties of resuspended mineral dust samples of defined size distributions were measured. Extinction coefficients (b_{ext}) and mass specific extinction cross sections (σ_{ext}) were determined for Saharan dust samples from different locations. The results for σ_{ext} were not very sensitive to the type of dust and varied at $\lambda=550$ nm between 3.3 ± 0.4 m² g⁻¹ and 3.7 ± 0.4 m² g⁻¹. The absorption coefficients (b_{abs}) and mass specific absorption cross sections (σ_{abs}) were determined with a novel multi-wavelength photo-acoustic absorption spectrometer (PAS). The single scattering albedo was close to 1 (0.98 to 0.99) at 532 nm and 1064 nm, but significantly lower (0.63 to 0.76) at 266 nm. Additionally the chemical and mineralogical composition of the dust samples were analysed with special regard to the iron oxide phases hematite and goethite. At $\lambda=266$ nm the mineral dust sample without any detectable iron oxides showed a significantly higher SSA compared to the sample with a hematite content of 0.6 wt-%.

We measured the specific extinction and absorption cross-sections of different airborne Saharan dust samples in the UV/VIS spectral region. All optical instruments sampled from a laboratory aerosol chamber which contained the mineral dust aerosol with a well characterised size distribution in the sub-micron particle diameter range.

Spectra of the extinction coefficient were measured from $\lambda=230$ to 1000 nm. The absorption coefficient, and hence the single scattering albedo were determined at three wavelengths in the near-UV ($\lambda=266$ nm), the visible ($\lambda=532$ nm), and the near-IR ($\lambda=1064$ nm). In addition to the optical measurements, the chemical and mineralogical compositions of the dust samples were analysed, with a focus on iron oxide mineral phases, i.e. hematite (Fe₂O₃) and goethite (FeO(OH)). The correlation between the deduced spectral quantities and the iron oxide content of the samples is discussed.

1 Introduction

Mineral dust aerosols from desert regions contribute significantly to the total atmospheric aerosol load (Andreae, 1995). During the transport over long distances the dust clouds are able to affect the earth's radiative budget by absorption and scattering in the near-UV, visible and infrared spectral regions. The optical properties vary with size and mineralogical composition of the dust, resulting in a positive or negative radiative forcing of the atmosphere. Especially dark mineral phases like iron oxides predominantly influence the wavelength-dependent mass specific absorption cross section of naturally occurring mineral dusts in the visible spectral region (Tegen et al., 1996; Sokolik and Toon, 1999).

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2 Aerosol chamber experiments

2.1 Experimental set-up and instrumentation

The experiments were conducted using the stainless steel aerosol chamber NAUA which is part of the AIDA (Aerosol Interactions and Dynamics in the Atmosphere) facility of Forschungszentrum Karlsruhe, Germany. The chamber has a volume of 3.7 m³ which allows measurements over several hours with the suite of devices used in this study (Fig. 1). The chamber is equipped with a mixing fan to achieve homogeneous conditions throughout the chamber volume. A constant synthetic air flow of about 2 m³ h⁻¹ balances for sampling losses and assures a well defined dilution rate of the aerosol during the experiment. All experiments were performed at constant temperature (296 K) and atmospheric pressure conditions.

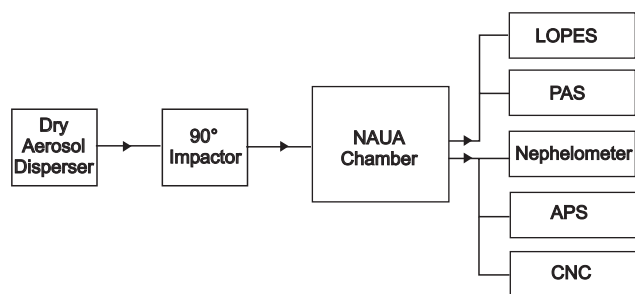


Fig. 1. Experimental setup of mineral dust dispersion, aerosol chamber NAUA and sampling. LOPES: Long Path Extinction Spectrometer, PAS: Photo-acoustic Absorption Spectrometer, APS: Aerodynamic Particle Sizer, CPC: Condensation Particle Counter.

Mineral dust aerosol was generated and added to the aerosol chamber by roughly dispersing the finest sieved fraction ($<20\ \mu\text{m}$) of the dry desert dust samples in a rotating brush disperser (Palas, RGB 1000), de-agglomerating the finest particles in an aerosol dispersion nozzle and removing larger particles in a series of four 90° inertial impactor stages. Both the brush disperser and the dispersion nozzle were operated with dry and particle-free synthetic air. The dispersion pressure of the nozzle was 1.5 bar. The impactor stages with a cut-off $d_{(50)}$ of $1.2\ \mu\text{m}$ (aerodynamic diameter) limited the size range of particles entering the NAUA chamber. Thus, particle losses by sedimentation in the horizontal sampling lines and the long horizontal flow tube of LOPES were minimised resulting in a low systematic error of less than 5% in the deduced specific optical cross sections. Figure 2 shows representative size distributions of the four mineral dust aerosols in the NAUA chamber during the experiments.

The extinction coefficient was measured wavelength-resolved from $\lambda=230$ to $1000\ \text{nm}$ with a spectral resolution of $2.5\ \text{nm}$ by the extinction spectrometer LOPES (Schnaiter et al., 2005). The optical cell of LOPES is a horizontally mounted stainless steel flow tube of $3.5\ \text{m}$ length and of $26\ \text{mm}$ inner diameter. The optical length is twice the geometric length. The tube was operated at a flow rate of $5\ \text{L}\ \text{min}^{-1}$ resulting in an aerosol transmission efficiency of more than 95% for particles with Stokes diameters smaller than $1\ \mu\text{m}$.

During calibration and validation the scattering coefficients at 450 , 550 and $700\ \text{nm}$ were measured for soot aerosol with a commercial integrating nephelometer (TSI; 3653). The instrument was calibrated by CO_2 and filtered air. The nephelometer was operated at a flow rate of $5\ \text{L}/\text{min}$.

From extinction and scattering measurements the absorption coefficients can be calculated by the difference method (DM):

$$b_{\text{abs}} = b_{\text{ext}} - b_{\text{sca}}$$

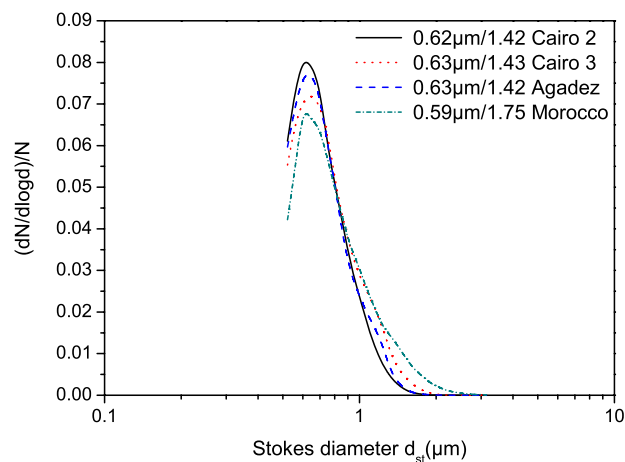


Fig. 2. Size distributions of mineral dust aerosols measured by the APS (Aerodynamic Particle Sizer). The aerodynamic equivalent particle diameter was converted to the Stokes diameter d_{st} using Eq. (1). The aerosol was sampled from the aerosol chamber NAUA.

Necessary nephelometer corrections are reasonable for moderately absorbing aerosols like biomass burning smoke with SSAs up to 0.7 (Schnaiter et al., 2005). For aerosols with higher SSA values the uncertainties in the calculated absorption coefficient increase due an increasing importance of the nephelometer and LOPES measurement errors. With regard to typical mineral dust aerosols with particle sizes above $0.5\ \mu\text{m}$, broad distributions and non-spherical particle shapes, the nephelometer turned out to be not applicable for precise measurements of the scattering coefficient, necessary to deduce the absorption coefficient by applying the DM. Closer examination of similar sized, non-absorbing quartz glass beads and non-spherical quartz glass particles indicates that the nephelometer correction depends not only on particle size but also on particle shape. Therefore, we used the DM only in case of highly absorbing soot to validate the direct absorption measurements by the photo-acoustic absorption spectrometer.

Absorption measurements were conducted with a multi-wavelength photo-acoustic absorption spectrometer (PAS) recently developed at the University of Szeged, Hungary, with a total sample rate of $1.5\ \text{L}\ \text{min}^{-1}$. PAS uses the fundamental ($\lambda=1064\ \text{nm}$), the frequency doubled ($\lambda=532\ \text{nm}$), and quadrupled ($\lambda=266\ \text{nm}$) modulated emission of a diode-laser-pumped, quasi-continuous wave, Q-switched Nd:YAG laser (Spectra Physics, laser head VHP80-106Q, power supply J20I-8S40-16NSI). The photo acoustic signal generation basically implies the conversion of light into acoustic energy. The intensity of the generated sound, which was detected with a microphone placed inside a resonant photo acoustic absorption cell (PA cell), is proportional to the concentration of absorbing trace molecules or aerosol particles. The Nd:YAG laser was run at $16\ \text{kHz}$ repetition rate (which

Table 1. Specific extinction cross sections, size distribution parameters.

sample	σ_{ext} ($\text{m}^2 \text{g}^{-1}$)			A	N	CMD	sd_g
	450 nm	550 nm	700 nm				
Cairo 2	3.6 ± 0.5	3.3 ± 0.4	2.8 ± 0.4	0.56 ± 0.1	6700	0.62	1.42
Cairo 3	3.9 ± 0.4	3.7 ± 0.4	3.2 ± 0.4	0.45 ± 0.1	1300	0.63	1.43
Agadez	4.0 ± 0.4	3.7 ± 0.4	3.2 ± 0.4	0.53 ± 0.1	1400	0.63	1.42
Morocco	3.8 ± 0.4	3.4 ± 0.4	2.9 ± 0.3	0.58 ± 0.1	2200	0.59	1.75

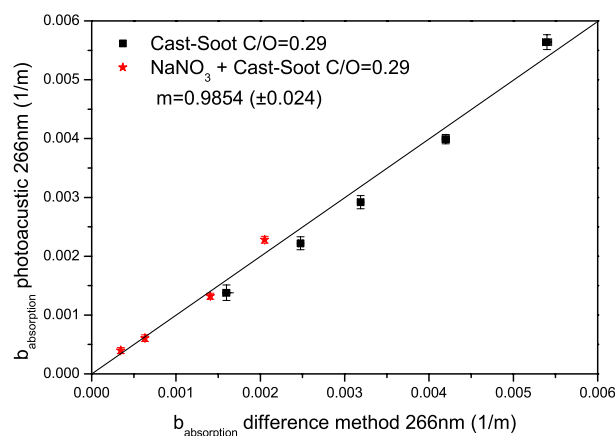
σ_{ext} – specific extinction cross section; A – Angström Exponent for extinction; N – initial number concentration at $t=0$; CMD – count median diameter; sd_g – geometric standard deviation

was the optimum frequency for the highest fourth harmonic power) and was modulate at 4 kHz (which was equal to the resonance frequency of the PA cells). The wavelength conversion was achieved by focusing first the near-infrared light (1064 nm) to a nonlinear LBO (LiB_3O_5) crystal followed by focusing the derived green light to a second nonlinear DKDP (deuterated potassium dihydrogen phosphate) crystal. The power of the laser lights were 750 mW at 1064 nm, 80 mW at 532 nm and 4 mW at 266 nm. This system is an upgraded version of a system used for photo acoustic ozone detection (Veres et al., 2005). While the ozone detection system used only one photoacoustic cell, for one wavelength of 266 nm, the present system uses 3 identical PA cells for the 3 available wavelengths. The aerosol sample from the chamber was split in three identical flows through the PA cells.

At $\lambda=266$ nm and $\lambda=532$ nm, the photo acoustic cells were calibrated with ozone and NO_2 , respectively. For the ozone calibration at $\lambda=266$ nm, a cross section of $9.44 \times 10^{-18} \text{ cm}^2/\text{molecule}$ was applied, taken from a spectrum by Voigt et al. (2001) measured at 293 K and 1000 mbar. For the NO_2 calibration at $\lambda=532$ nm, a cross section of $1.47 \times 10^{-19} \text{ cm}^2/\text{molecule}$ was used, taken from a spectrum by Voigt et al. (2002) which was also measured at 293 K and 1000 mbar.

At the IR-wavelength $\lambda=1064$ nm, the PAS system was calibrated with a polydisperse soot aerosol from a co-flow diffusion flame of propane and air (Combustion Aerosol Standard, CAST, Jing-CAST Technologies). The burner was operated at a C/O ratio of 0.29. Wavelength-dependent mass-specific absorption cross sections of the same CAST-soot in the wavelength-range between 450 nm and 700 nm have been determined in a separate campaign, yielding an Angström exponent for absorption of 1.0 (Schnaiter et al., 2006). Taking this Angström exponent the absorption at $\lambda=1064$ nm was calculated by extrapolation. The uncertainty of calibration was determined to be $\pm 2\%$ for all three wavelengths.

Correlations between absorption coefficients measured with PAS and derived from extinction and scattering measurements by the difference method (DM, Schnaiter et al., 2005) are shown in Figs. 3 and 4 for pure CAST soot and mixtures of CAST soot with non-absorbing but scattering

**Fig. 3.** Comparison of PAS at 266 nm with the difference method (DM) using CAST-soot with a C/O ratio of 0.29.

sodium nitrate at $\lambda=266$ nm and $\lambda=532$ nm, respectively. The latter experiments with external aerosol mixtures showed that the presence of a strongly scattering aerosol fraction does not interfere significantly with the PA measurement of absorption. Here, the absorption Angström exponent determined by DM in the 450 to 700 nm spectral region was used to extrapolate and interpolate the DM deduced absorption coefficients at $\lambda=450$ nm and $\lambda=550$ nm to the PAS wavelengths at $\lambda=266$ nm and $\lambda=532$ nm. The detection limit of the PAS system was determined to be 10^{-5} m^{-1} . The overall uncertainty of the PAS measurement is about 5% for our aerosol concentrations. Total particle number concentrations of the aerosols in the chamber were measured with a condensation particle counter (TSI, CPC 3022A) operated at a flow rate of 0.3 L min^{-1} . Number size distributions were measured with an aerodynamic particle sizer (APS 3321, TSI) operated at a flow rate of 5 L min^{-1} . The APS acquires size distributions in the aerodynamic diameter range from 0.5 to $10 \mu\text{m}$ with a resolution of 32 channels per decade.

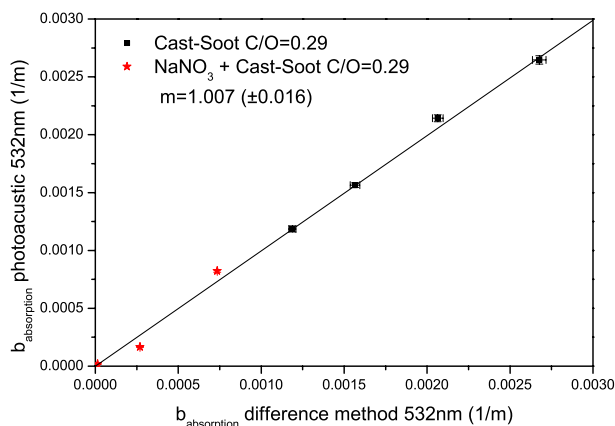


Fig. 4. Comparison of PAS at 532 nm with the difference method (DM) using CAST-soot with a C/O ratio of 0.29.

2.2 Experimental procedure

At the beginning of each experiment the NAUA chamber was evacuated, flushed, and re-filled with dry synthetic air, resulting in a background particle concentration less than 2 cm^{-3} . All experiments were performed at room temperature and ambient pressure. Mineral dust aerosol was then added to the chamber using the dispersion set-up described above for about 30 min, which led to initial particle number concentrations between 1300 cm^{-3} and 6700 cm^{-3} . The container volume served as an aerosol reservoir for the sampling instruments. On the time scale of the experiments, aerosol dilution was the dominant loss process for mineral dust particles which had no significant effect on the specific optical properties of the aerosol. In addition to the online measurements mineral dust samples were collected on Nuclepore filters for analysis by scanning electron microscopy (SEM).

Two sets of experiments were performed: In a first set, the optical properties were quantified by means of extinction measurements solely, together with measurements of the aerosol number concentration and size distribution. In a second set of experiments, the PAS instrument was also operated at the NAUA chamber to measure optical absorption coefficients simultaneously to the extinction coefficients.

3 Aerosol concentration and size distribution

A series of similar experiments was performed with four Saharan mineral dust samples of different colour and origin. A soil sample from the Northern Sahara, Cairo 2, was collected 70 km northeast of Cairo city from a depth of 0.5 m. The colour of this sample is light yellow to light brown. A wind blown dust sample, Cairo 3, was collected on a flat surface in Cairo city after a dust storm event at 19 March 2002. This sample has a similar colour as Cairo 2. A Sahelian soil sam-

ple from Agadez (Niger) and a soil sample from Morocco became available by courtesy of Lothar Schütz from the University of Mainz. Compared to the Cairo samples, these samples have a darker appearance and are coloured yellow brown (Agadez) to reddish brown (Morocco).

All samples were sieved to grain size fractions of less than $20 \mu\text{m}$, 20 to $75 \mu\text{m}$, and greater than $75 \mu\text{m}$. Our experiments were performed with the fraction of particles less than $20 \mu\text{m}$. The other fractions were disregarded for the present study.

Applying the same dispersion procedure resulted in a larger aerosol number concentration for the Cairo 2 sample (6700 cm^{-3}) compared to the other samples (1300 to 2200 cm^{-3}). This indicates an enhanced number fraction of particles smaller than $1.2 \mu\text{m}$ aerodynamic diameter in the Cairo 2 sample (Table 1).

Since the SEM images revealed a non-spherical, irregular, but compact shape of the dust particles (Figs. 5 to 8), the aerosol mass concentration was determined from the APS measurements by first calculating the volume equivalent sphere diameter d_{st} (Stokes diameter) of the irregularly shaped mineral dust particles

$$d_{st} = \sqrt{\frac{\chi}{\rho_p}} d_{ae} \quad (1)$$

from the measured aerodynamic diameter d_{ae} . The particle density $\rho_p = 2.5 \text{ g cm}^{-3}$ was determined by pycnometric analysis of the mineral dust sample. To account for the irregular particle shape we used a dynamic shape factor $\chi = 1.5$ which is typical for minerals like quartz (Baron and Willeke, 2001). Figure 2 shows a representative size distribution of our four dust samples as function of d_{st} . A lognormal size distribution was then fitted to the resulting aerosol size distributions. Good representations of the measured size distributions were achieved by this approach especially for the larger particles which represent the main contribution to the total aerosol volume. The moments of the fit result, i.e. the count median diameter (CMD), the geometric standard deviation, and the total particle number concentration, were used to calculate the mass concentration of the aerosol taking the particle density given above. During an experiment both the CMD and geometric standard deviation remained almost constant. For the interpretation of the measured aerosol optical properties it is important to note that comparable CMDs and standard deviations were found for the four investigated mineral dust types (Table 1).

4 Mineralogical composition and analysis

The desert dust samples were characterised in cooperation with the Institute for Materials Research (IMF I) at FZK, the division of Water Technology and Geotechnology, Institute for Technical Chemistry (ITC-WGT) of FZK and the Astrophysical Institute of the University of Jena.

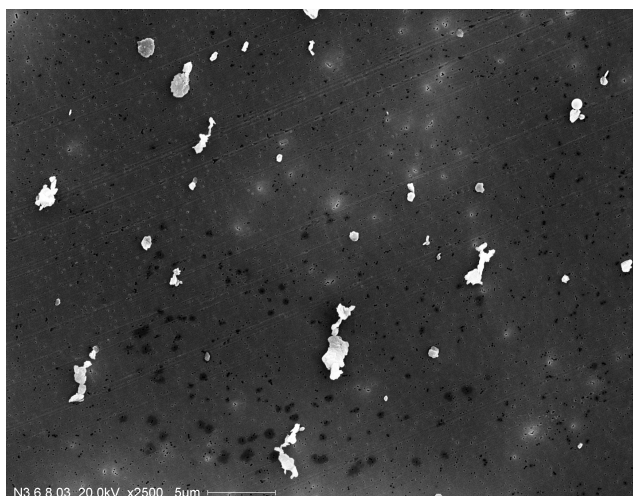


Fig. 5. SEM picture of mineral dust Agadez.

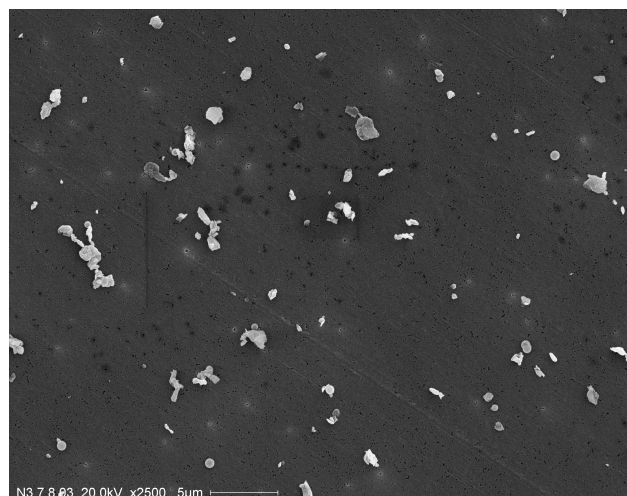


Fig. 7. SEM picture of mineral dust Cairo 3.

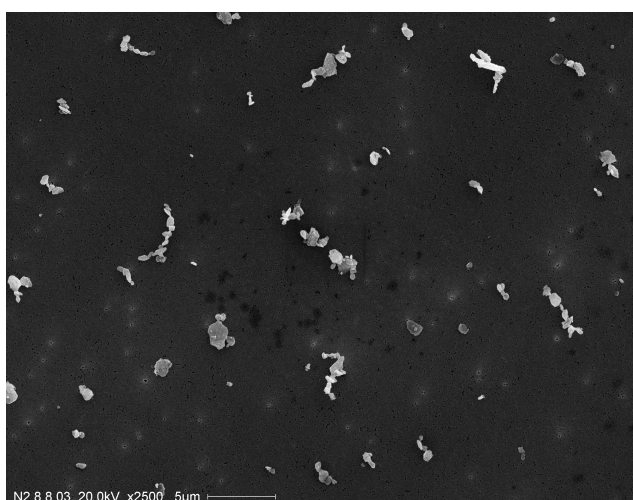


Fig. 6. SEM picture of mineral dust Cairo 2.

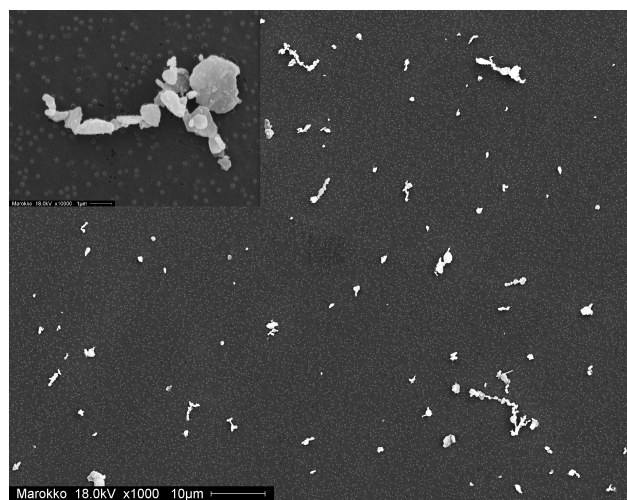


Fig. 8. SEM picture of mineral dust Morocco (note the lower magnification compared to the other pictures). The inserted magnified image gives an impression of the diversity of the dust grains in terms of size and surface structure.

The elemental composition of each mineral sample was determined for the dust fraction $<20\ \mu\text{m}$ by X-Ray Fluorescence Analysis (XRF, Bruker AXS, SRS 303AS). The samples were ignited for one hour at 1000°C prior to the analysis. For each dust type a sample of 200 mg was analyzed. For the measured elements weight percentages of oxides were calculated (Table 3). Thus, the samples yielded total iron oxide content, which mainly contributes to the absorption coefficient of the dust aerosol, and ranged between 2.0 to 5.0 wt-%.

In addition to the chemical analysis the mineral composition was investigated by IR spectroscopy and X-ray-powder-diffractometry with Rietveld phase analysis (XRD, Bruker AXS, D8 diffractometer). For IR spectroscopy the dust fraction $<20\ \mu\text{m}$ was mixed with KBr powder before the mixture was pressed to a pellet. The pellet was then analysed by

IR transmission spectroscopy in the $50\ \text{cm}^{-1}$ to $2000\ \text{cm}^{-1}$ spectral range. The XRD analysis was performed directly on a granular sample of the dust fraction $<20\ \mu\text{m}$. Figure 9 shows the IR-spectra of Cairo 2 and Agadez in comparison to the spectrum of the pure clay mineral Montmorillonite. Corresponding absorption bands of sample- and reference-spectra at 500 and $1000\ \text{cm}^{-1}$ indicate the presence of related clay minerals in both Saharan samples. In Fig. 10 the spectra of Cairo 2 and Agadez are compared with the spectrum of the pure calcite. While Cairo 2 shows several absorption bands consistent with the spectrum of calcite at $1500\ \text{cm}^{-1}$, $800\ \text{cm}^{-1}$ and below $500\ \text{cm}^{-1}$ the Agadez sample displays no significant absorption bands at these wavenumbers.

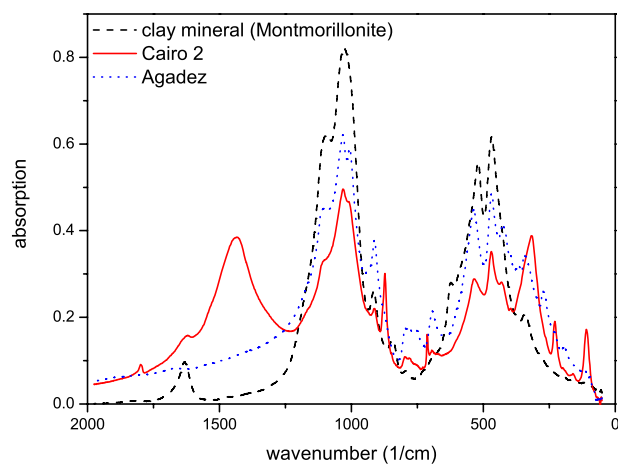


Fig. 9. IR-spectra of mineral dust Cairo 2, Agadez and of the clay mineral Montmorillonite.

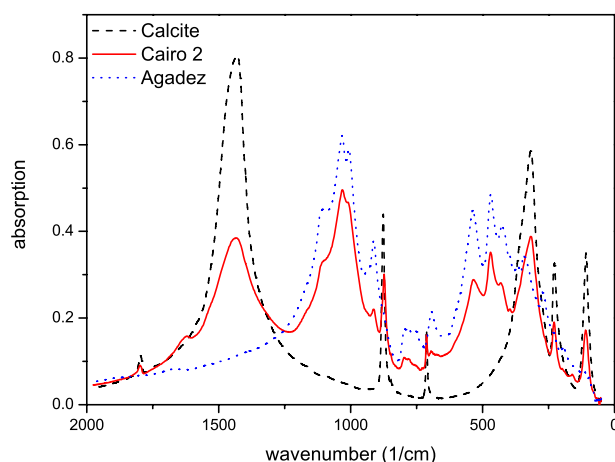


Fig. 10. IR-spectra of mineral dust Cairo 2, Agadez and of calcite.

Both methods, IR spectroscopy and XRD analysis, show that all investigated samples contain quartz, feldspars and clay minerals in different compositions. In addition, the samples from Cairo and Morocco contain significant amounts of calcite, dolomite and gypsum.

In our case the XRD method was applied to identify the iron oxide phases hematite or goethite. Depending on the crystal phase of interest the method has a detection limit of 0.1 to 0.5 wt-% for iron oxides. While the Cairo 2 and Cairo 3 samples showed no detectable iron oxide phases, there are indications for goethite in the Agadez sample in a very low concentration of ~ 0.2 wt-% which is at the detection limit of the method.

Only for the Morocco sample the XRD method yielded, after careful analysis, a detectable amount of 0.6 ± 0.1 wt-% of hematite. However, due to a possible enrichment of hematite in the fine aerosol fraction, the hematite concentration in the chamber aerosol could be significantly higher than the concentrations found in the XRD analysis of the granular samples.

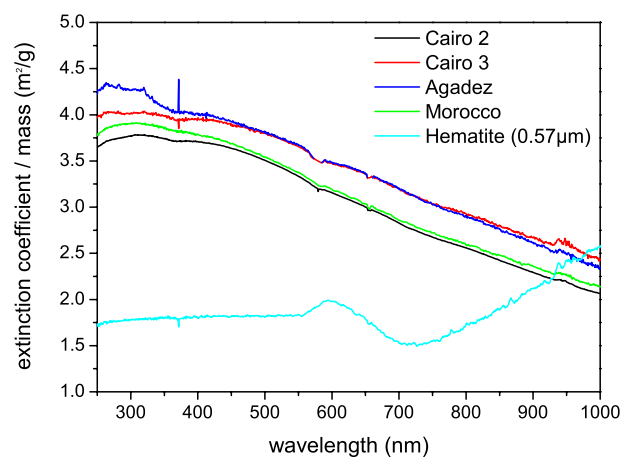


Fig. 11. Extinction spectra of the investigated mineral dust samples and hematite.

5 Results

UV-VIS extinction spectra are shown in Fig. 11 for the four mineral dusts and pure hematite. As expected from the wavelength-dependent complex refractive index of pure hematite (Sokolik and Toon, 1999) the extinction spectrum of synthetic hematite aerosol differs significantly from that of the investigated natural mineral dusts. The hematite sample consists of a narrow size distribution of pseudo-cubic hematite particles (Sugimoto et al., 1993). Probably the hematite spectrum shows a Mie-type resonance structure around 600 nm, but this does not affect the general absorption-induced flat spectral behaviour in the wavelength range below 600 nm, which is obviously a specific spectral feature of hematite.

For calculating σ_{ext} the aerosol mass concentration was derived from the measured size distributions. The total uncertainty in the mass determination of 20% resulted from errors of total number concentration, density and shape factor of the dust. By error propagation these uncertainties were considered in the calculation of the specific optical cross sections given in Table 1 and Table 2.

Table 1 gives σ_{ext} at $\lambda=450$, 550 and 700 nm and the corresponding Angström exponent of extinction in the 450 to 700 nm spectral region. It is obvious from Table 1 and Fig. 11 that σ_{ext} and its wavelength-dependence is not very sensitive to the type of dust. Since in case of weakly absorbing mineral dust aerosol the extinction is dominated by scattering, this result reflects the comparable size distributions of the different dust types in the chamber (Table 1). The wavelength-dependent specific absorption cross-section σ_{abs} of the mineral dust aerosols were deduced from the PAS measurements. The resulting values are given in Table 2 together with the corresponding values of the single scattering albedo (SSA) deduced from the PAS absorption and

Table 2. Specific absorption cross sections and single scattering albedos.

sample	$\sigma_{\text{abs}} \text{ m}^2 \text{ g}^{-1}$			SSA		
	266 nm	532 nm	1064 nm	266 nm	532 nm	1064 nm
Cairo 2	0.81±0.16	0.02±0.004	0.03±0.005	0.76±0.02	0.99±0.001	0.99±0.01
Morocco	1.10±0.26	0.06±0.014		0.63±0.04	0.98±0.002	

SSA – single scattering albedo; σ_{abs} – specific absorption cross section

Table 3. Element analysis of desert sand samples by X-ray fluorescence analysis calculated as oxides (sd = standard deviation).

Elements as	Cairo 2		Cairo 3		Agadez		Morocco	
	wt-%	sd	wt-%	sd	wt-%	sd	wt-%	sd
SiO ₂	55.2	0.08	62.0	0.06	79.4	0.31	70.22	0.17
Al ₂ O ₃	7.7	0.01	9.4	0.01	11.5	0.06	8.37	0.01
CaO	24.3	0.05	13.9	0.01	0.7	0.01	11.47	0.02
MgO	2.9	0.03	4.7	0.04	0.4	0.00	2.26	0.01
Fe ₂ O ₃	4.5	0.02	5.0	0.02	2.0	0.01	3.63	0.03
K ₂ O	1.18	0.003	1.7	0.00	3.6	0.01	2.04	0.01
Na ₂ O	1.1	0.18	1.1	0.01	2.2	0.13	0.52	0.02
P ₂ O ₅	0.2	0.01	0.2	0.00	0.0	0.01	<0.1	
TiO ₂	1.4	0.07	1.2	0.02	0.3	0.01	0.81	0.05
Loss of ignition	19.1		16.8		2.7		12.1	

the LOPES extinction coefficients. The PAS instrument was only available for measurements with the light-coloured dust sample Cairo 2 and the reddish dust sample from Morocco. At the visible wavelength of $\lambda=532$ nm SSAs of 0.99 ± 0.001 and 0.98 ± 0.002 were deduced for the Cairo 2 and Morocco samples. In contrast to these relatively high SSA values in the visible, quite low values of 0.76 ± 0.02 and 0.63 ± 0.04 were found for the Cairo 2 and Morocco sample, respectively in the near-UV at $\lambda=266$ nm (Table 2). This strong wavelength dependence of the SSA reflects the strong wavelength dependence of the absorption coefficient resulting in Angström exponents as large as 5.3 and 4.2 for the Cairo 2 and Morocco samples, respectively. When comparing absorption Angström exponents of mineral dusts which are based on a few sampling wavelengths, one has to keep in mind that the corresponding absorption spectra might have significant band structures (e.g. in case of hematite), which is in contrast to the absorption spectra of soot aerosols. For Cairo 2 σ_{abs} increases from $0.02\pm 0.004 \text{ m}^2 \text{ g}^{-1}$ at $\lambda=532$ nm to $0.81\pm 0.16 \text{ m}^2 \text{ g}^{-1}$ at $\lambda=266$ nm and for Morocco from $0.06\pm 0.014 \text{ m}^2 \text{ g}^{-1}$ at $\lambda=532$ nm to $1.1\pm 0.26 \text{ m}^2 \text{ g}^{-1}$ at $\lambda=266$ nm. Additionally for Cairo 2 σ_{abs} was determined at the IR-wavelength $\lambda=1064$ nm to be $0.03\pm 0.005 \text{ m}^2 \text{ g}^{-1}$, with a SSA of 0.99 ± 0.01 .

6 Discussion

Although the elemental analysis yielded total iron oxide contents between 2.0 and 5.0 wt-% for all mineral dust samples studied in this work, cf. Table 3, neither hematite nor goethite could be detected by XRD in the Cairo 2 and Cairo 3 samples. In case of the Agadez sample goethite is probably present in a low concentration of about 0.2 wt-%. The Morocco sample is the only sample which contained a detectable hematite mass fraction of 0.6 ± 0.1 wt-%. Differences between the quantified total iron content detected by XRF (Table 3) and the hematite or goethite contents determined by XRD could be due to difficulties of the Rietveld method associated with poor crystallographic ordering in mineral dusts. Discrepancies may also be attributed to different sample preparation, i.e. mechanically versus manually ground samples (Weidler et al., 1998), and extraction procedures.

While Alfaro et al. (2004) found in their dust samples comparable total iron contents in the range of 3.0 to 6.5 wt-%, they assigned the significant amount of 2.8 to 5.8 wt-% of this iron to be present in iron oxide phases. However their iron oxide content was determined with the method by Lafon et al. (2004) as the difference between total iron and its structural fraction. Structural fraction in this case quantifies the remaining iron content after applying an adapted reductive extraction method as commonly used in soil science (Lafon

et al., 2004). This method provides no structural information about the extractable iron and therefore cannot distinguish between the presence of goethite or hematite in the samples.

For calculating SSAs and specific absorption cross sections Alfaro et al. (2004) determined scattering and absorption coefficients by an integrating nephelometer and a dual wavelength aethalometer. Due to light scattering by particles deposited in the filter matrix of the aethalometer the measured attenuation over-estimates the particle absorption and, therefore, has to be corrected. Alfaro et al. (2004) applied a correction method by Bond et al. (1999) with two experimentally determined correction terms, while using the uncorrected scattering cross sections of the nephelometer. They investigated one Sahelian dust from Niger, one Saharan dust from Tunisia and one dust from China and found SSAs at $\lambda=660$ nm of 0.95 ± 0.01 , 0.97 ± 0.01 and 0.97 ± 0.01 for these dusts. These values are slightly lower compared to our results (Sect. 5). However, since the SSA is strongly size-dependent (Tegen et al., 1996), this difference might be due to their bimodal dust size distributions with a second mode in the super-micron size range (2.0 to 5.0 μm) in addition to the sub-micron particle mode. The values for σ_{abs} of Alfaro et al. (2004) range between 0.12 and 0.07 $\text{m}^2 \text{g}^{-1}$ at 325 nm and between 0.02 and 0.01 $\text{m}^2 \text{g}^{-1}$ at 660 nm which is in the same order of magnitude as our results.

Retrievals of remote sensing data by Cattrall et al. (2003) from Saharan dust passages at Dry Tortugas, Florida revealed SSA values across the visible spectral range. At $\lambda=530$ nm the SSA specified by Cattrall et al. (2003) is about 0.96 ± 0.2 in good agreement with our results. By comparing their and others retrieval results with SSA model results based on laboratory data of the complex refractive indices of bulk minerals they found that the remotely sensed values result in a significantly weaker dust absorption than model predictions.

At the high-altitude aerosol station Jungfraujoch Coen et al. (2004) measured scattering coefficients with an integrating nephelometer and absorption coefficients with an aethalometer. During Saharan dust events (SDE) they found Angström exponents of absorption of 1.5 to 1.8 between $\lambda=370$ nm and $\lambda=590$ nm which were significantly larger than the usual Angström exponents of 1.0 to 1.2 from non-Saharan background aerosol. Reflections on the filter were corrected by an empirical correction factor, which was calculated for different aerosol types by Weingartner et al. (2003). As already mentioned in the previous section the determination of Angström exponents of absorption for our dust samples based on the PAS absorption measurements at $\lambda=532$ nm and $\lambda=266$ nm results in relatively high values between 4.2 (Morocco) and 5.3 (Cairo 2). These results indicate that there exists a significant wavelength-dependence in the UV-VIS light absorption by mineral dusts.

7 Conclusions

For the quantification of the spectral optical properties of weakly absorbing aerosols, in particular mineral dust, the presented combination of extinction spectroscopy with LOPES and absorption measurements with the multiple-wavelength PAS proved to be very useful. For two mineral aerosols we found SSA values which are comparable to values observed in recent laboratory and field studies. Moreover, the SSA values determined in the present study show a strong wavelength-dependence with a steep decrease from the visible to the near-UV.

The chemical and mineralogical composition of the dust samples were analysed with special regard to the iron oxide phases hematite and goethite. Although we have measured the absorption coefficients of only two mineral dust samples, these first investigations indicate that the presence of iron oxide phases in the dust significantly increases the absorption cross section of the mineral dust aerosol in the visible and especially in the UV. Further laboratory investigations with pure hematite and clay minerals as well as with external mixtures of both will reveal the dependence of σ_{abs} and SSA on iron oxide content more clearly.

The comparison of optical measurements to the chemical/mineralogical analysis implies investigations on the similar grain fractions. Therefore sufficient sample for the chemical and mineralogical analysis from the resuspended fraction have to be obtained. Besides the chemical and mineralogical analysis presented here, we will focus more on single particle analysis by SEM/EDX in future investigations. It will also be important to address in more detail the issue of mineralogical composition as function of the particle size.

Additionally, the PAS instrument will be improved by including another wavelength in the near-UV at $\lambda=355$ nm which allows a more reliable determination of the Angström exponent and a better comparison of our absorption data with other laboratory and field results.

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